

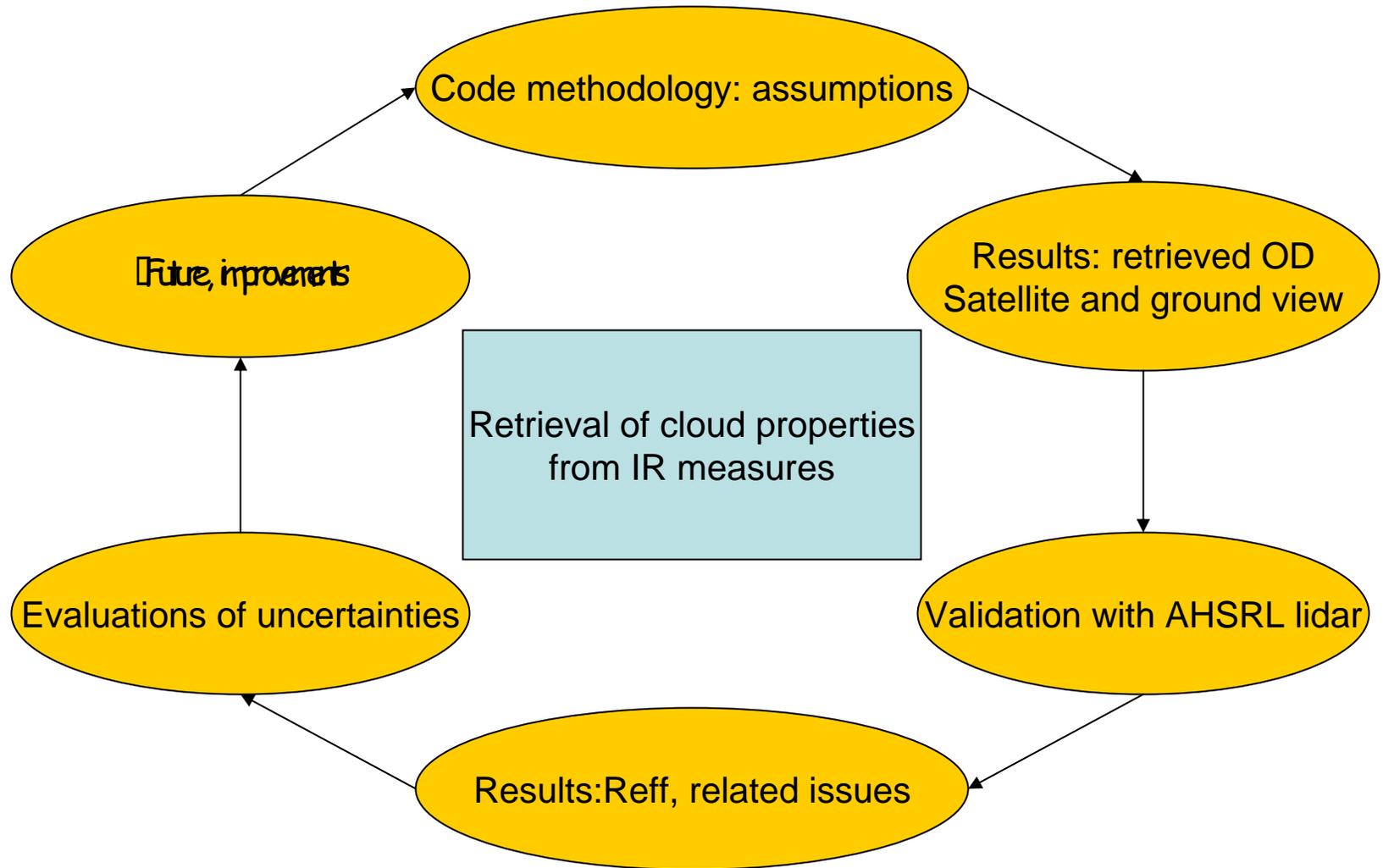
A NEW CLOUD PROPERTIES RETRIEVAL METHODOLOGY FOR IR MEASUREMENTS:

An application to MPACE 2004 field experiment

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OVERVIEW:



Goal:

Use the best information available to derive a validation product for selected geographical sites
Satellite applications (?)

RT-RET: the Retrieval scheme/1

ORIGIN: Modification of a LbL-MS code based on doubling and adding algorithm (Evans and Stephens; Rizzi and Miskolczi)

A-PRIORI INFORMATION:

1. Cloud geometrical boundaries (→ lidar/CO₂ slicing/...)
2. T and wv profiles (→ ECMWF re-analysis grib files/radiosondes)

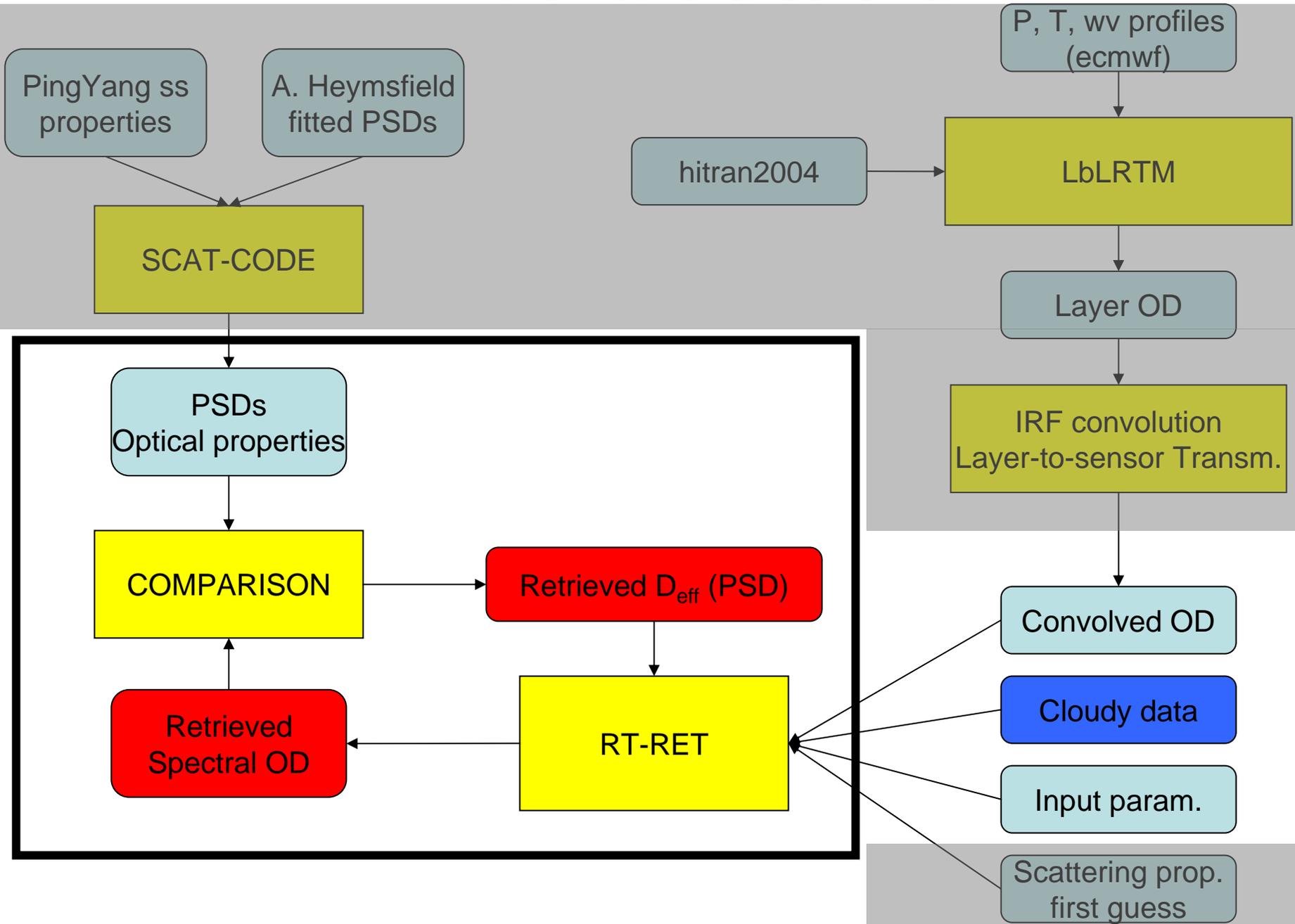
ASSUMPTION:

1. Cloud OD is homogeneous in the layer (investigate!)
2. Plane parallel geometry and multiples layer cloud

NEW FEATURES:

1. Optical depths space
2. Simplified phase function but multiple scattering processes accounted for
3. Applicable to multiple hyper-spectral sensors (ground-airborne-satellite)

RT-RET: the Retrieval scheme/2



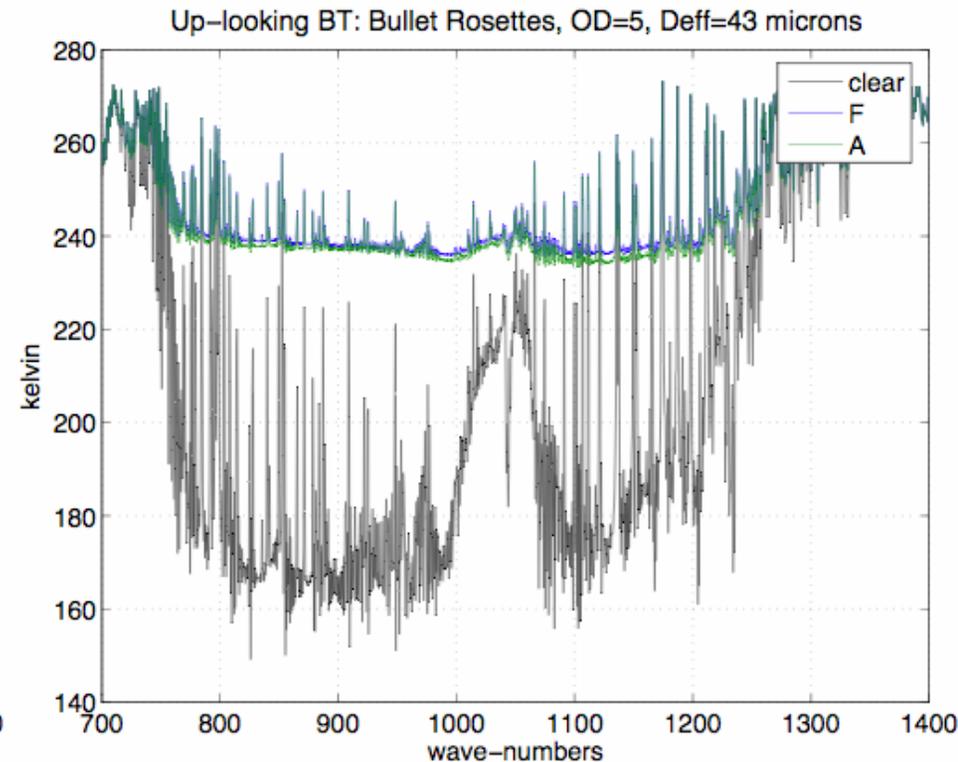
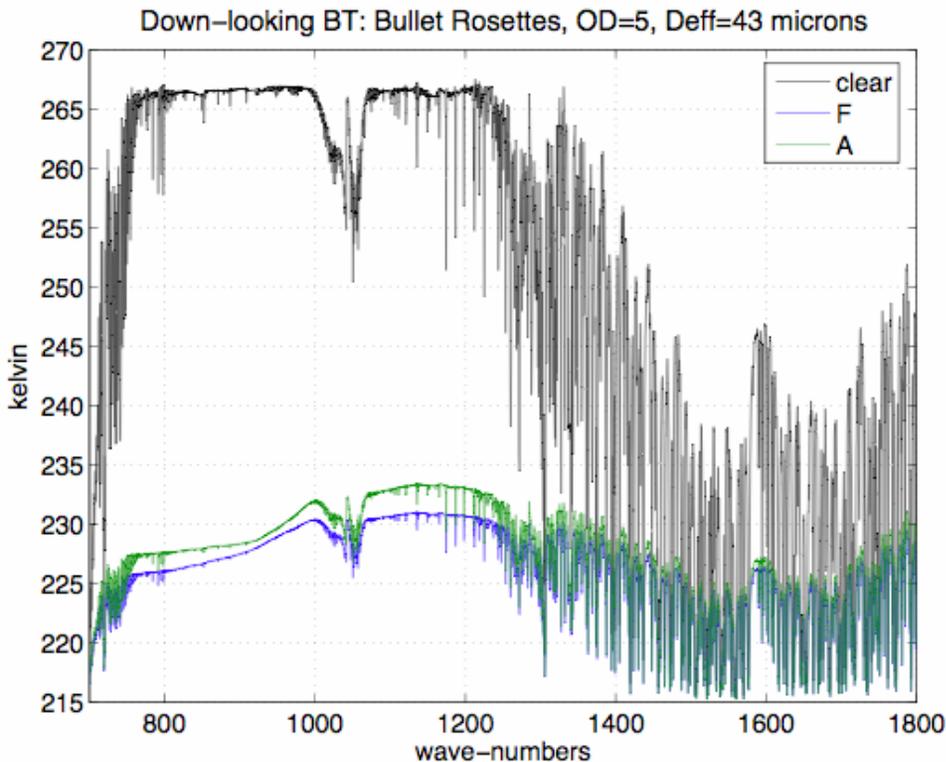
RT-RET: the Retrieval scheme/3

Scattering is a second term correction in the infrared window both from above and either below cloud measurements. Nevertheless it cannot be neglected

Clear solution -----

F=full solution -----

A=absorption only -----

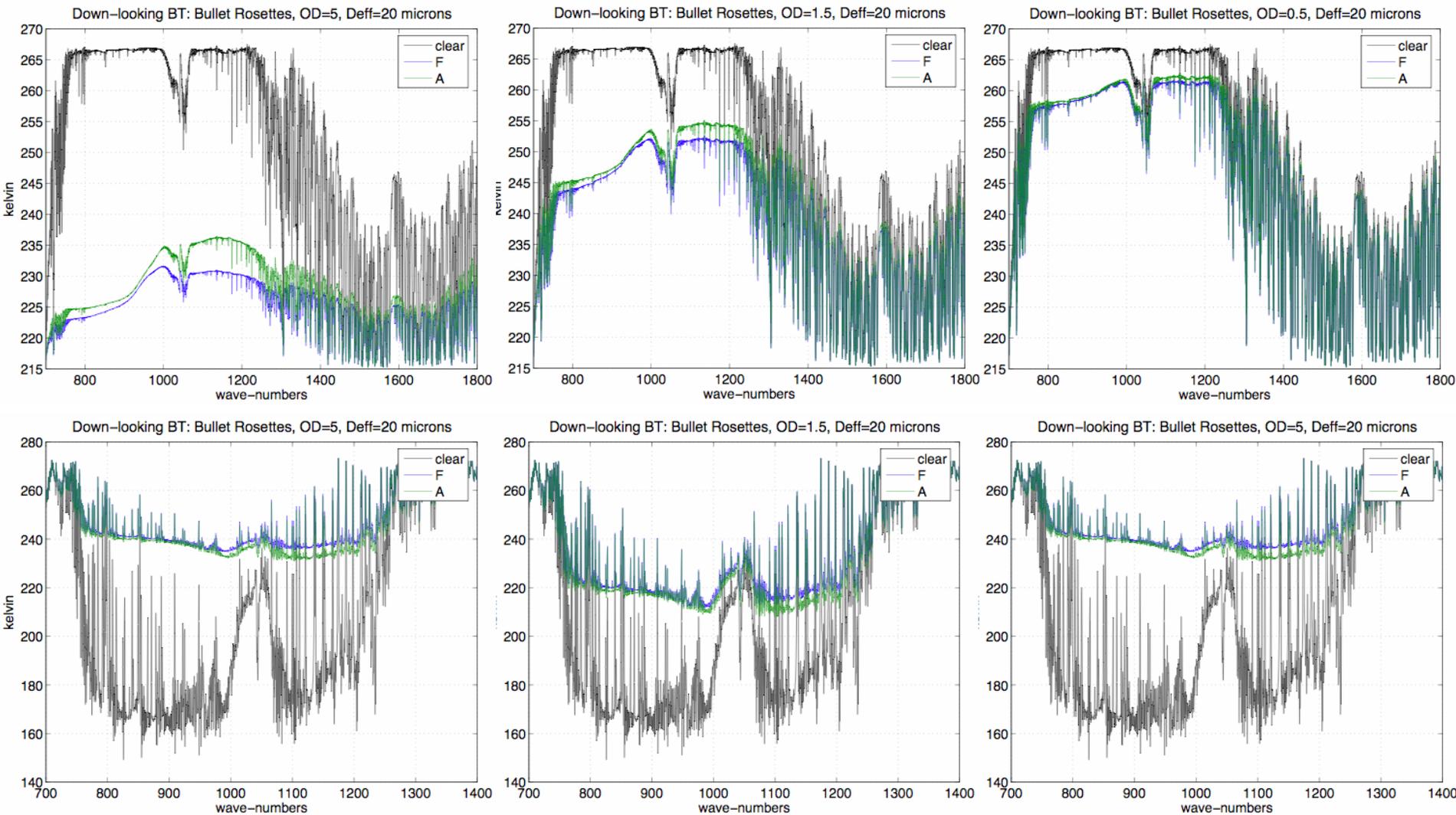


Knowledge of the T and wv profile is extremely important especially when up-looking.

RT-RET: the Retrieval scheme/4

The spectral shape of the radiance in the 800-1000 cm^{-1} window is mainly driven by the absorption processes independently of the size of the ice crystals

The slope of the radiance (esp. from above) in the 800-1000 cm^{-1} window band is not only habit dependent, but also OD and surface emissivity dependent (boundary conditions are very important)

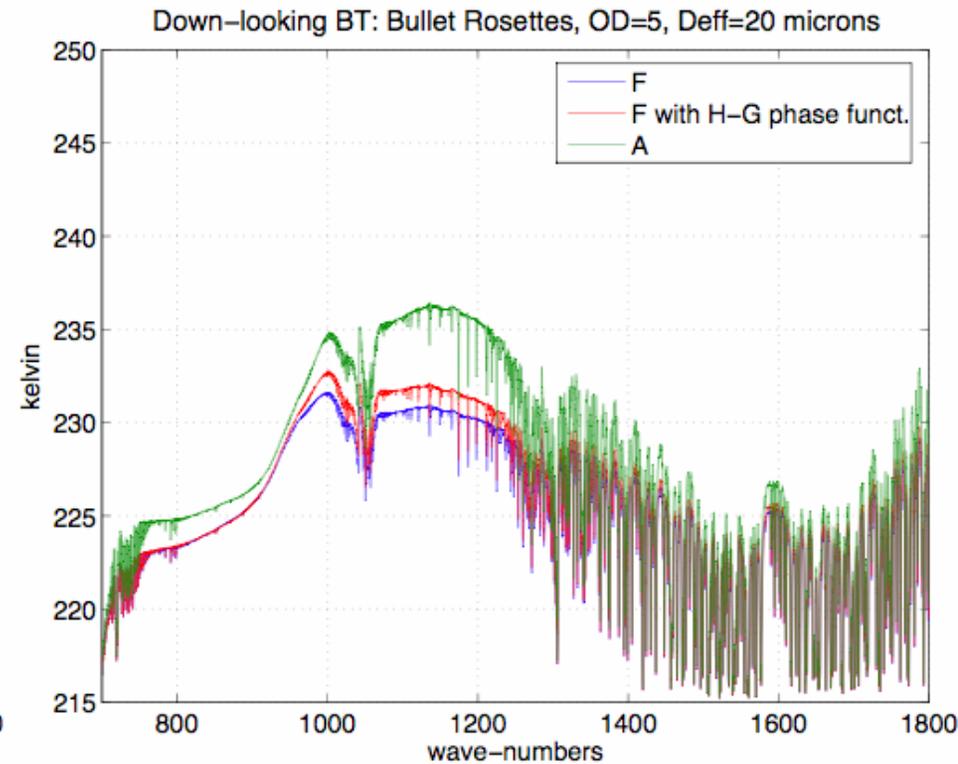
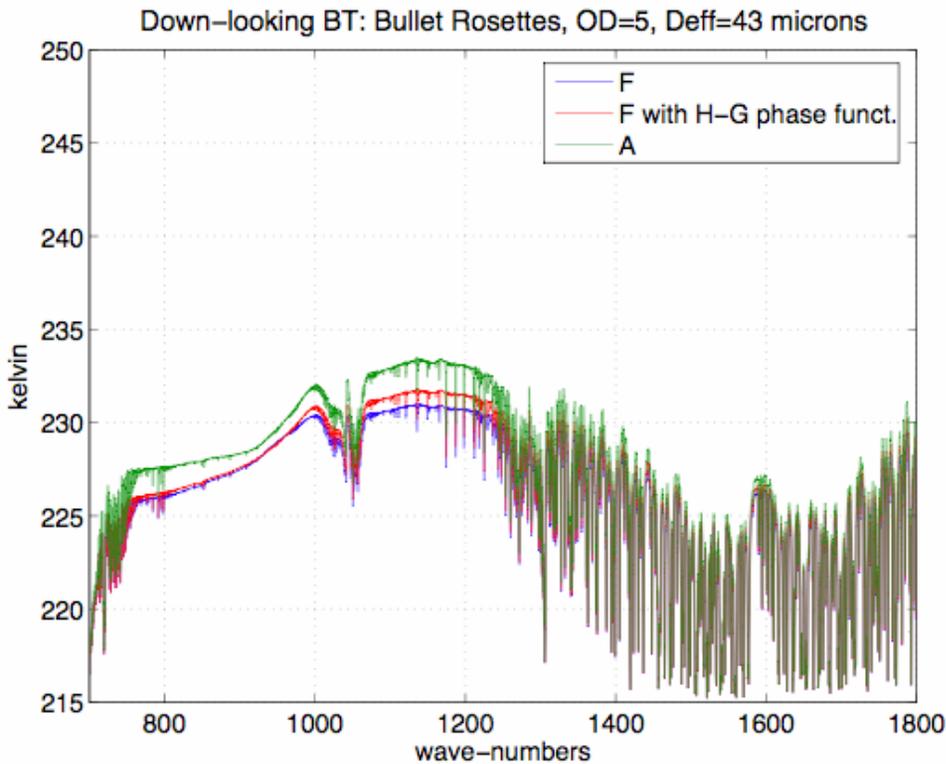


RT-RET: the Retrieval scheme/5

In the 800-1000 cm^{-1} window band the scattering properties of the PSDs are weaker dependent on the kind of habit (or mixtures of habits) constituting the PSD with respect to the 1050-1250 cm^{-1} band that contains more habit's information

Moreover a simplified phase function appears to represent a good approximation in that band.

Heeney-Greenstein phase function correction.



RT-RET: the Retrieval scheme/6

For every wave-number RT-RET finds the “fake” OD that makes the forward model computation match the data. The first guess scattering properties are obtained using spectral: ω and g . The precision is user dependent and usually set to a threshold less than 0.5% in the radiance values.

Wave-numbers are analyzed **sequentially**.

To speed up computations the last OD value is used as input for the next wave-number analyzed

Convergence to the real value is obtained by changing the OD as follows:

$$\Delta OD = \frac{\Delta R}{(F_s - F_c) \cdot t_c}$$

From the assumption that the up-ward radiance is: $R \approx t_c \cdot (\varepsilon_s F_s) + (1 - t_c) \cdot F_c$ and that the only variables are $t_c(OD)$ and R

$$\Delta OD = \frac{\Delta R}{(-F_c) \cdot t_c}$$

From the assumption that the down-ward radiance is: $R \approx (1 - t_c) \cdot F_c$ and that the only variables are $t_c(OD)$ and R

Where:

t_c = cloud transmissivity

R = measured radiance

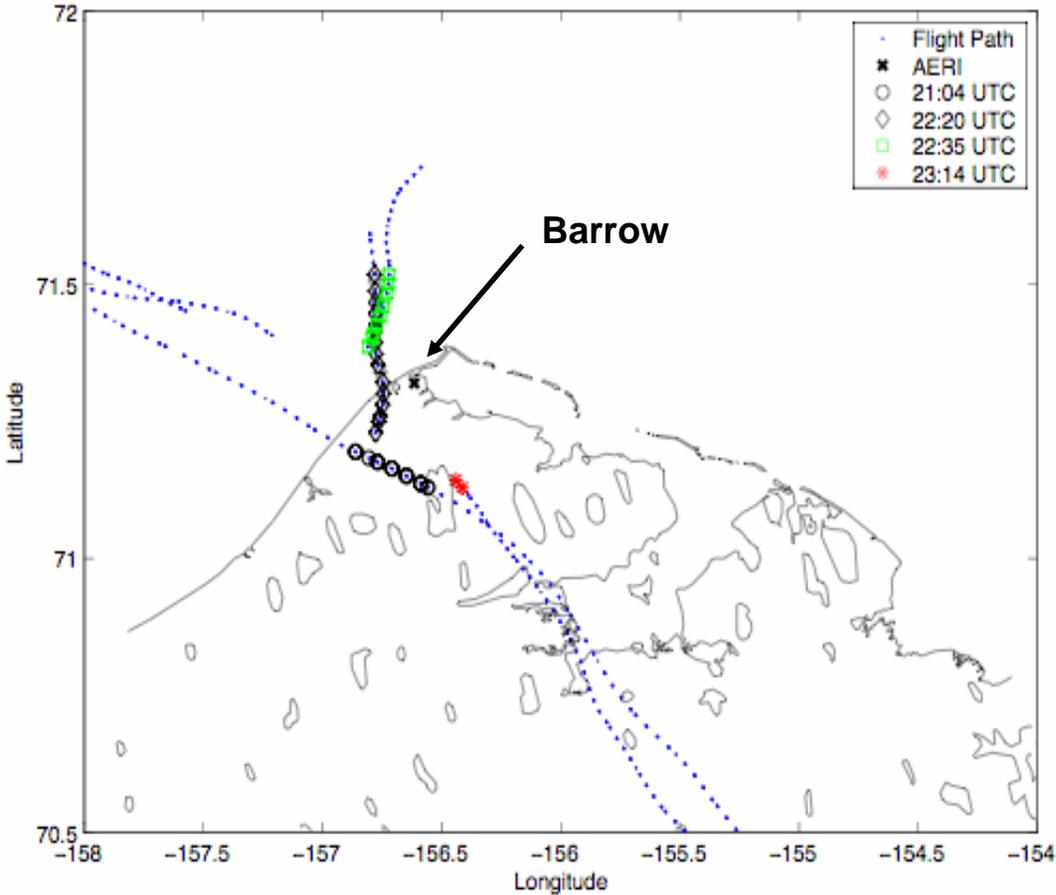
F_c = cloud radiance

F_s = surface radiance

ε_s = surface emissivity (assumed 1)

MPACE: October 17th 2004

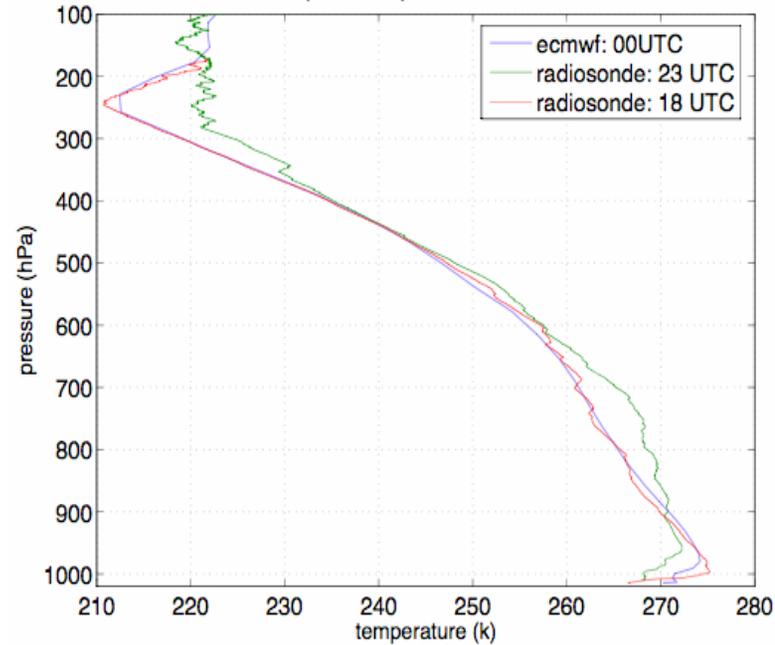
Proteus Flight Track (MPACE) 17 Oct 2005 over Barrow ARM Site



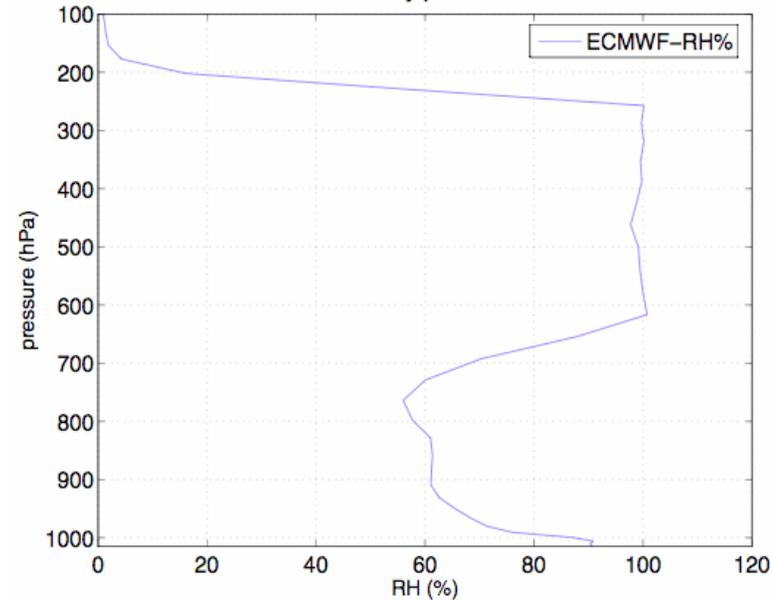
QuickTime™ and a
TIFF (LZW) decompressor
are needed to see this picture.

Hydrometeor from
Cloud Particle
Imager, CPI on
Citation (Verlinde et
al.)

Temperature profile: 2004/10/18



Relative Humidity profile: 2004/10/18



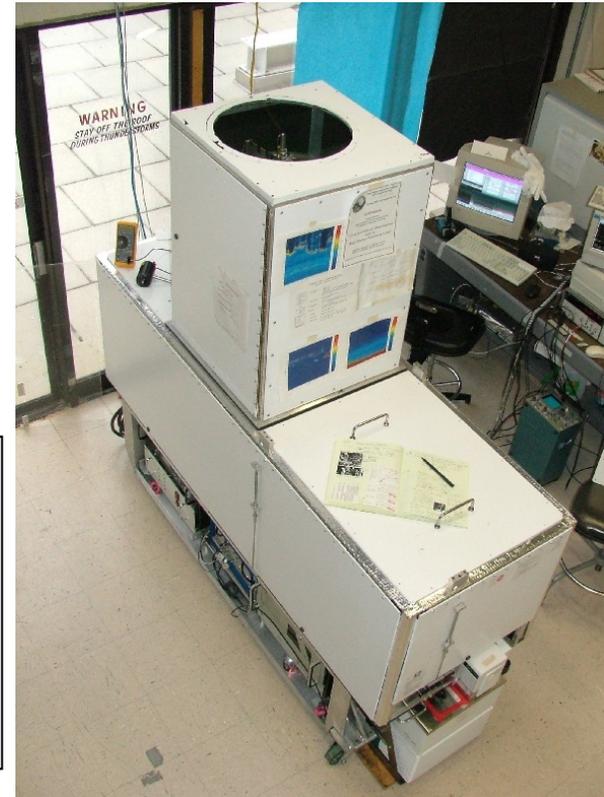
AHSRL

The Arctic High Spectral Resolution Lidar

Multi-channel lidar capable of independent measurements of the cloud depolarization, extinction, and backscatter cross-section. Two signals can be processed to yield separate lidar returns from aerosol and molecular scattering. Separation is possible because the wavelength spectrum of the molecular lidar return is Doppler broadened by molecular thermal motion. The separation of molecular and aerosol returns permits the HSRL to measure the extinction and aerosol backscatter cross-sections independently.

Laser Wavelength:	532 nm
Length of laser pulse:	40 ns
Pulse repetition rate:	4 kHz
Laser power:	~0.4 W
Receiver FOV:	45 mrad
Altitude Resolution:	7.5 m
Maximum Recorded Alt:	30 km

The AHSRL specifications



Thanks to Ed Eloranta (SSEC)

IR Sensors

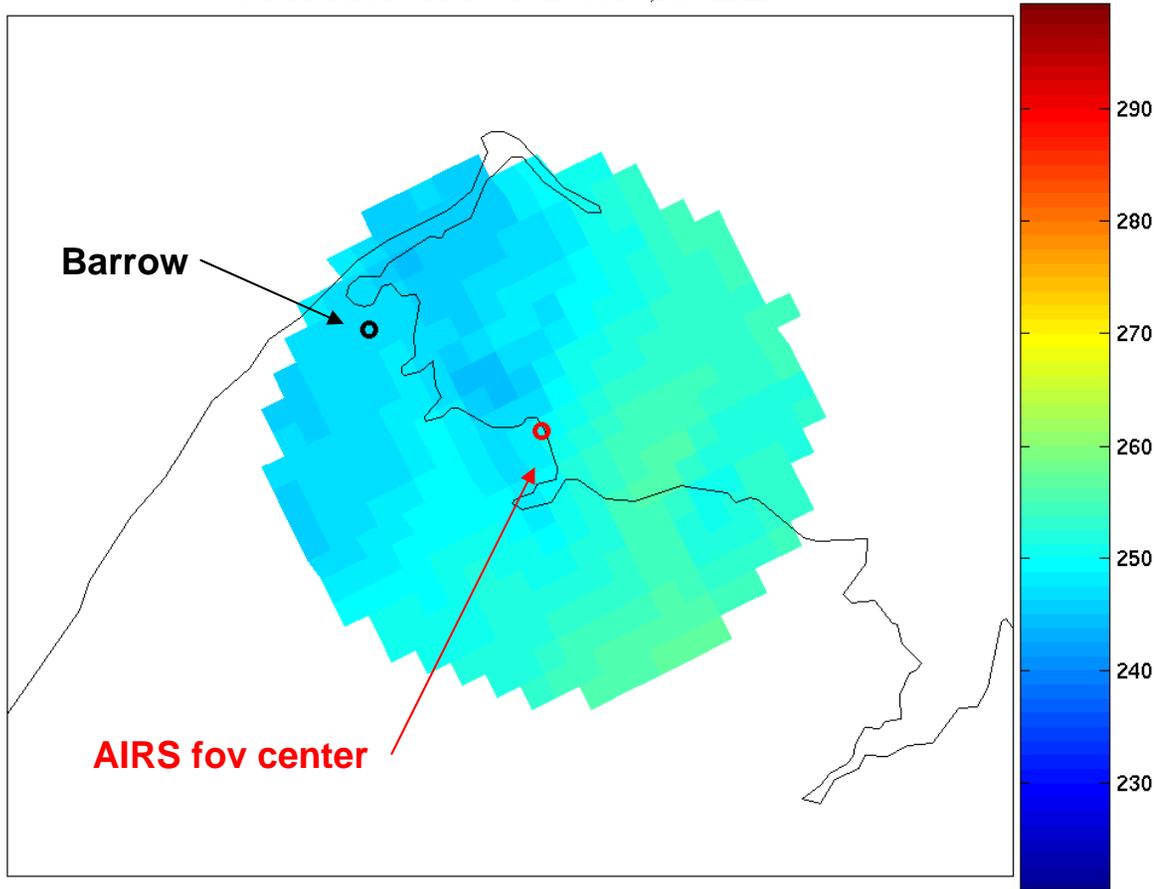
AIRS: Atmospheric Infrared Sounder. Satellite (AQUA)

S-HIS: Scanning High-resolution Interferometer Sounder. Airborne (on Proteus)

AERI: Atmospheric Emitted Radiance Interferometer (in rapid scan). Ground-base (Barrow)

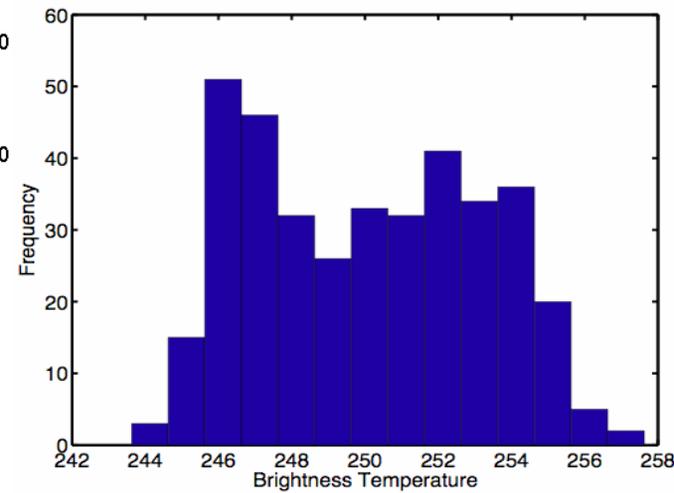
AIRS data: Modis collocation

AIRS Collocated MODIS 11um BT Oct. 17, 2004 22:22:04



Advection of mid-upper level cloud

Modis BT Histogram

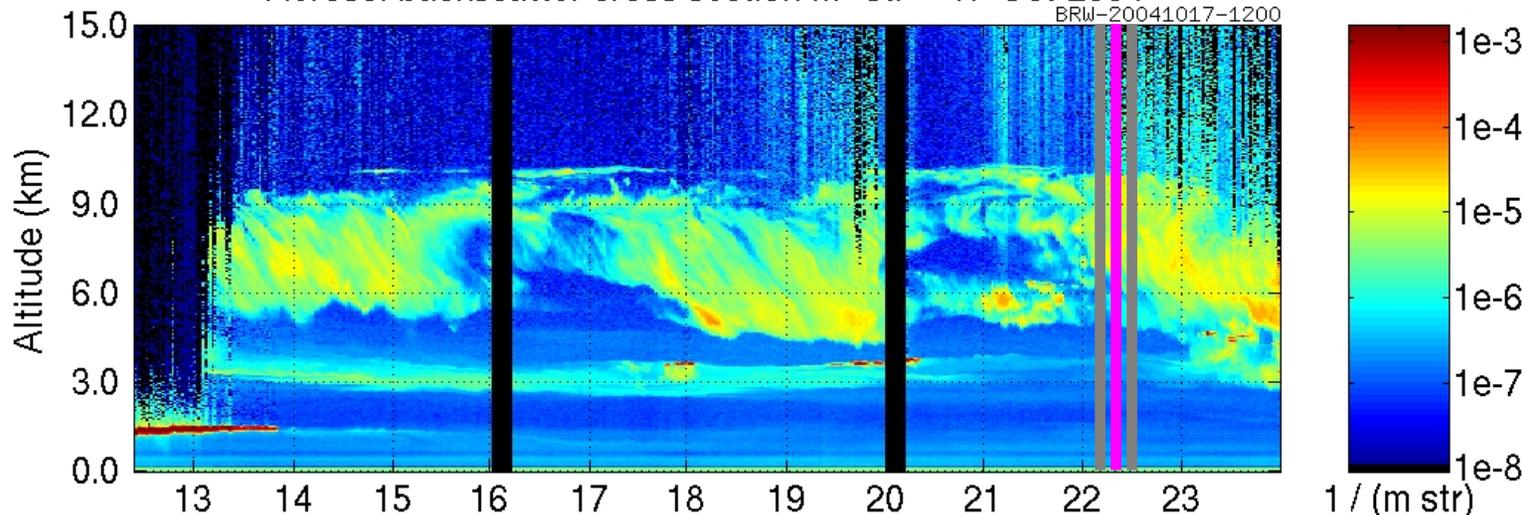


AHSRL data

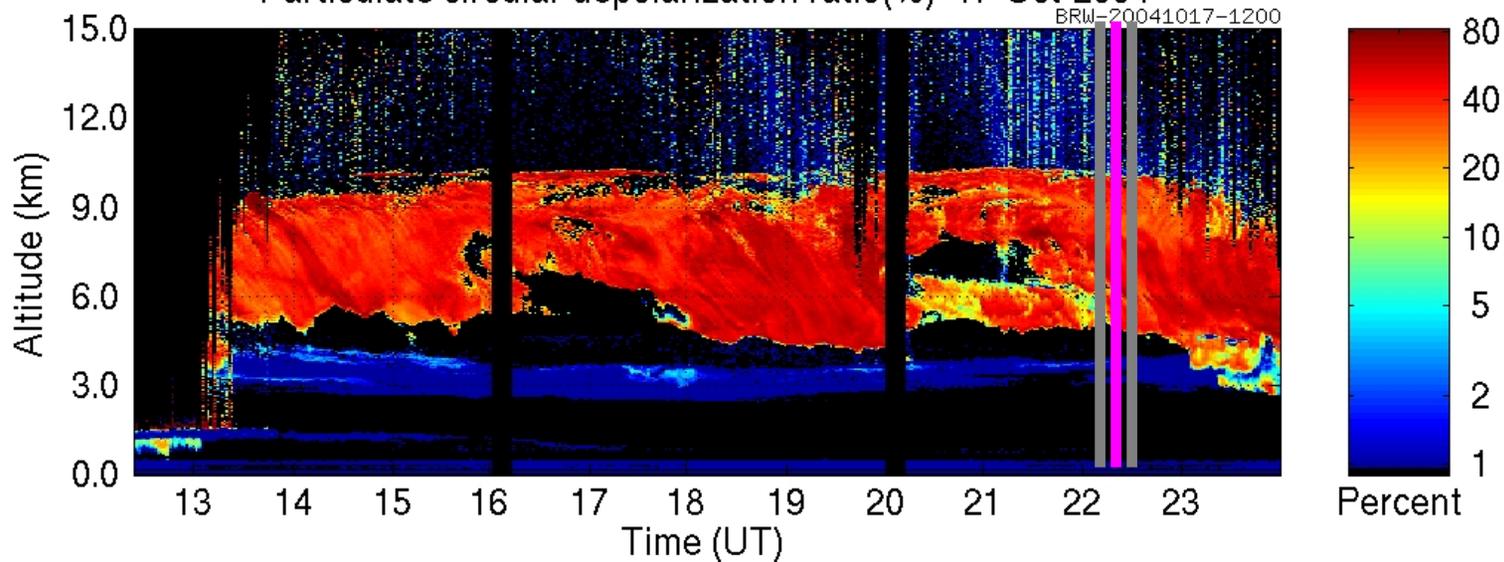
AIRS overpass

S-HIS time coincident passes

Aerosol backscatter cross section $\text{m}^{-1}\text{str}^{-1}$ 17-Oct-2004

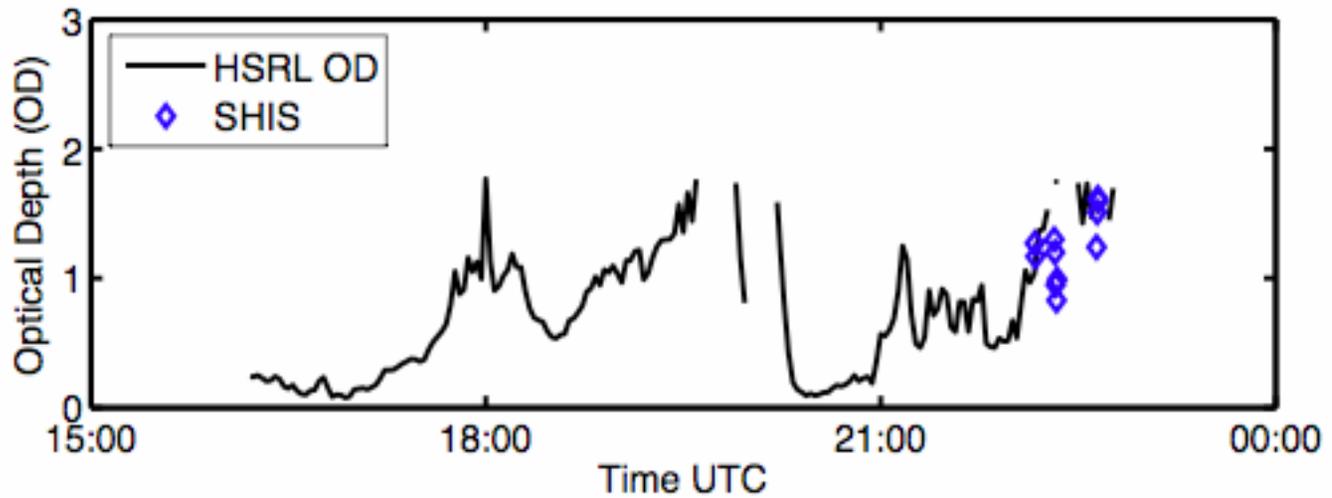
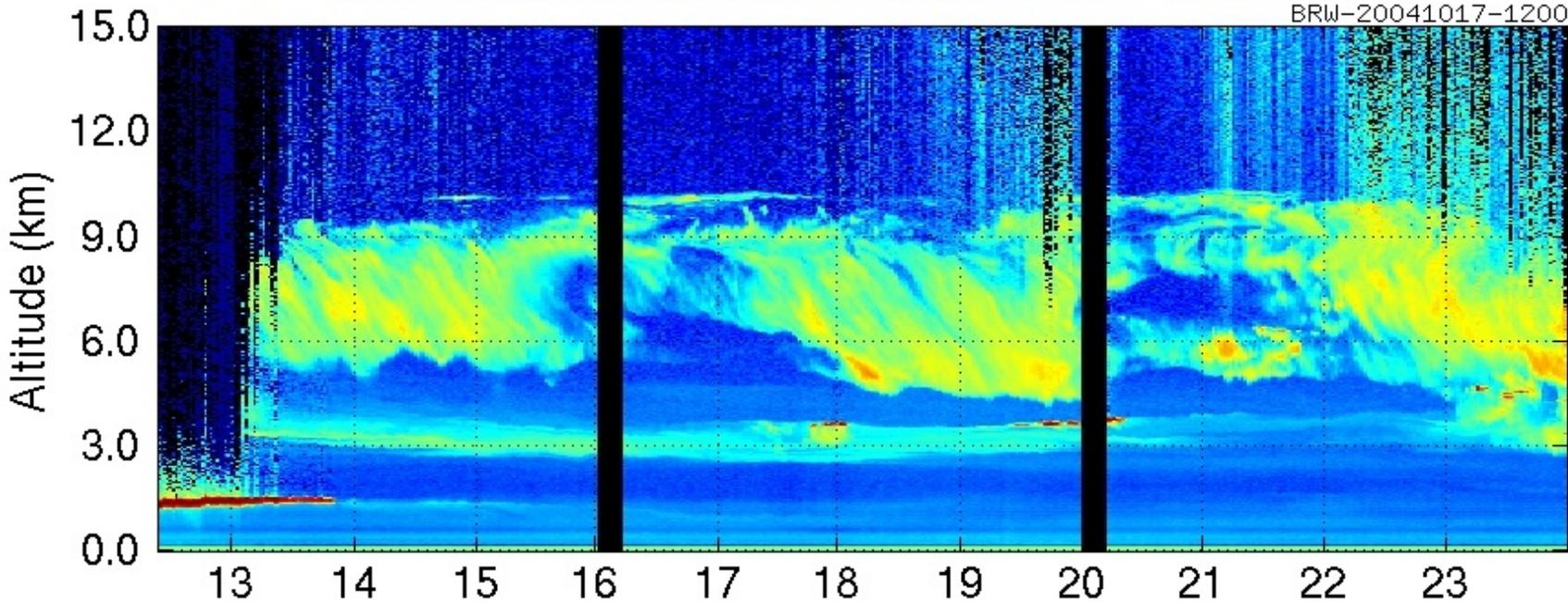


Particulate circular depolarization ratio(%) 17-Oct-2004



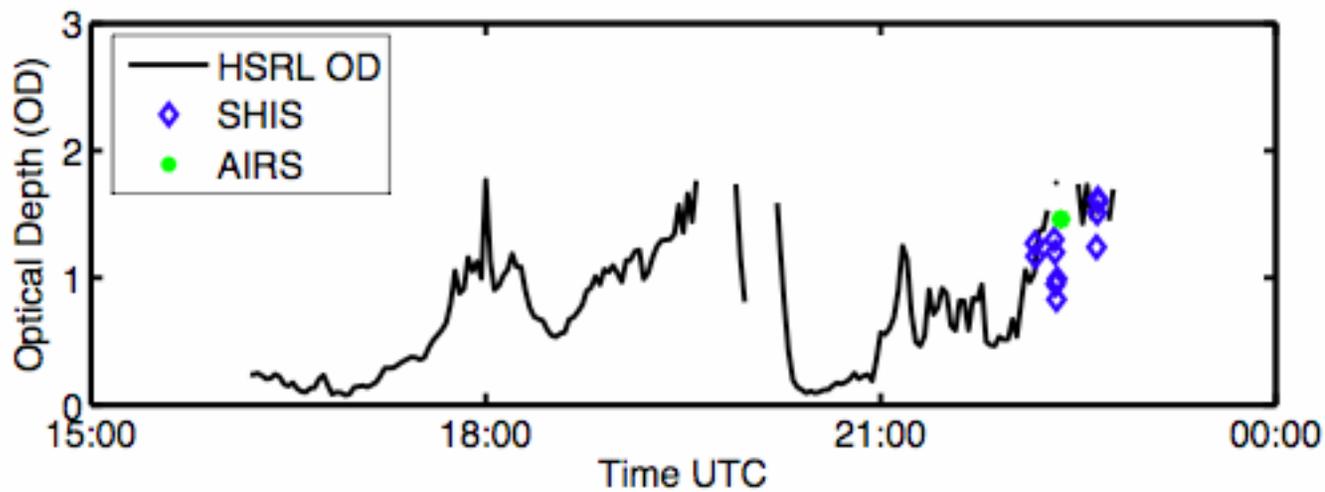
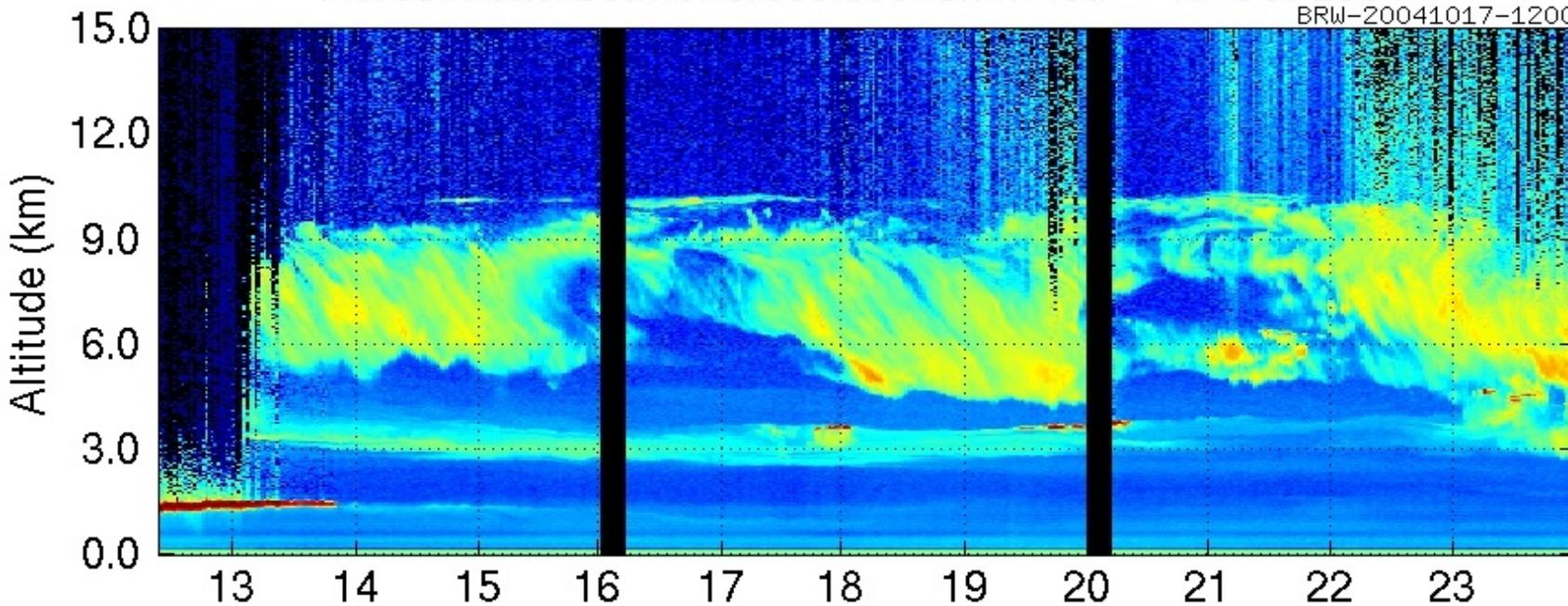
Aerosol backscatter cross section $m^{-1}str^{-1}$ 17-Oct-2004

BRW-20041017-1200



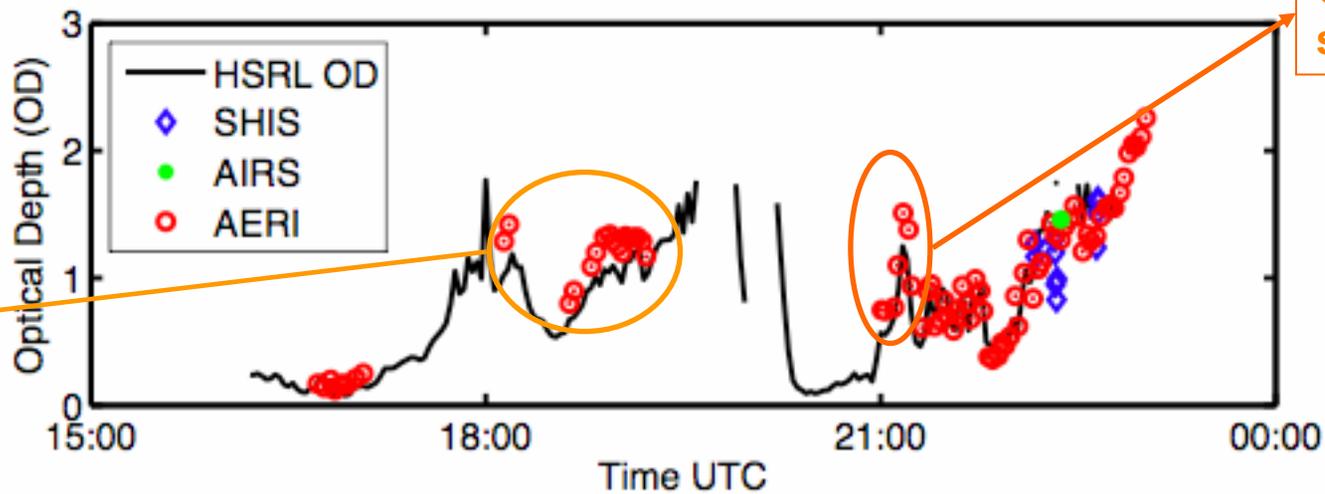
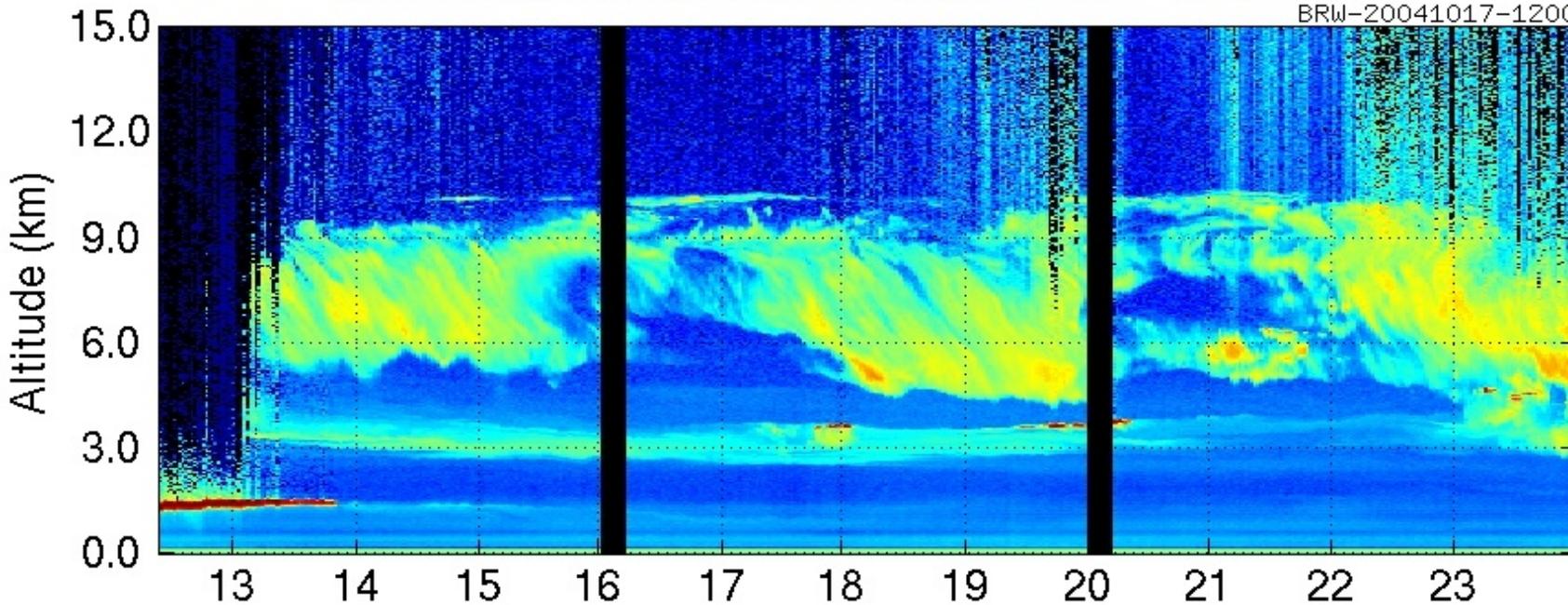
Aerosol backscatter cross section $\text{m}^{-1}\text{str}^{-1}$ 17-Oct-2004

BRW-20041017-1200



Aerosol backscatter cross section $m^{-1}str^{-1}$ 17-Oct-2004

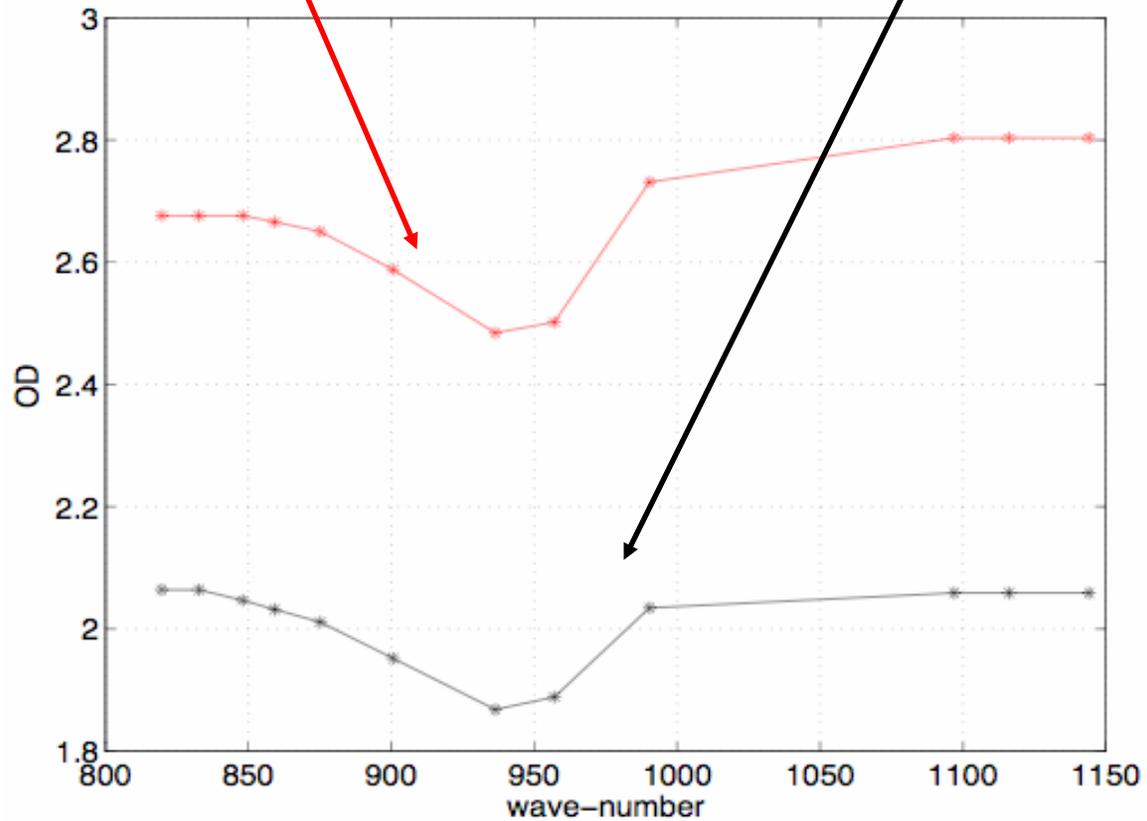
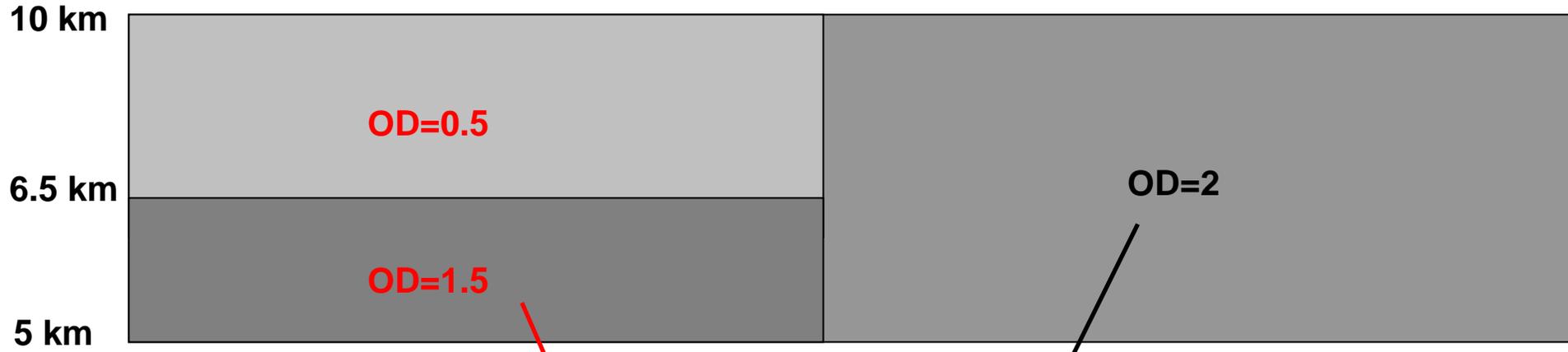
BRW-20041017-1200



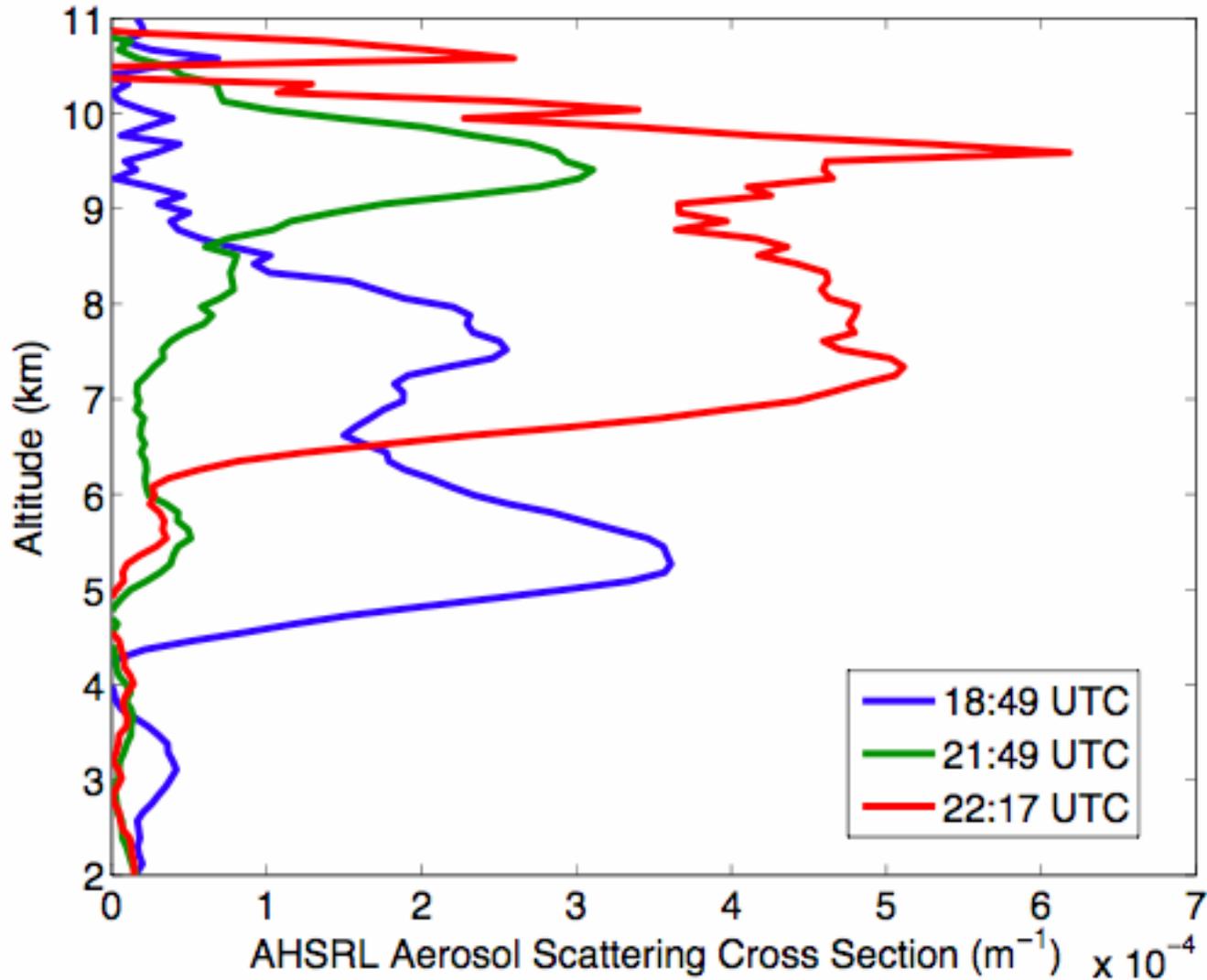
Water cloud and OD spread

OD spread

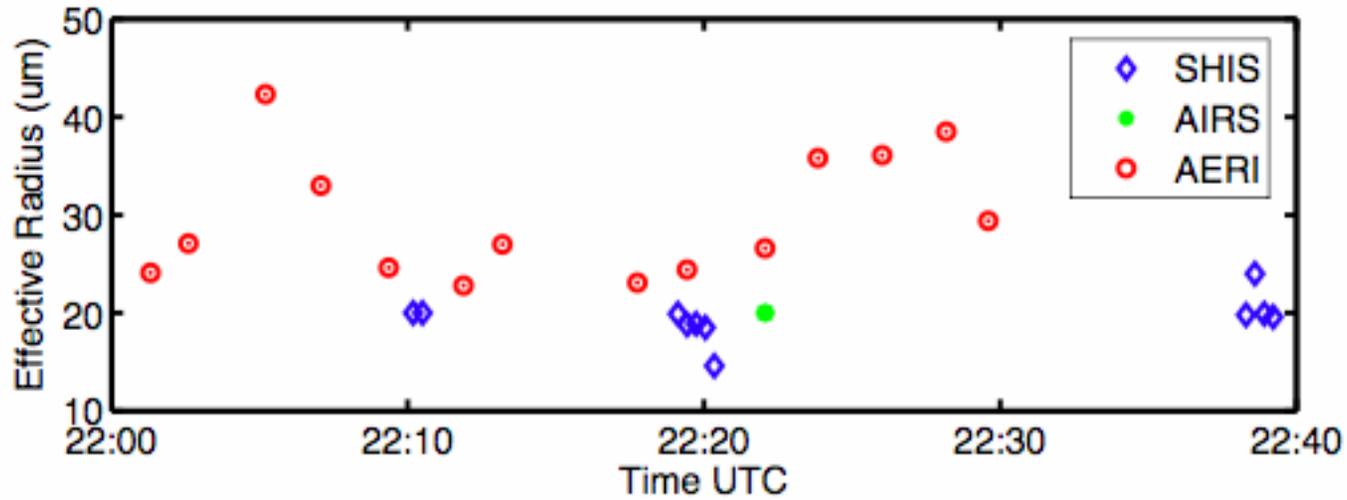
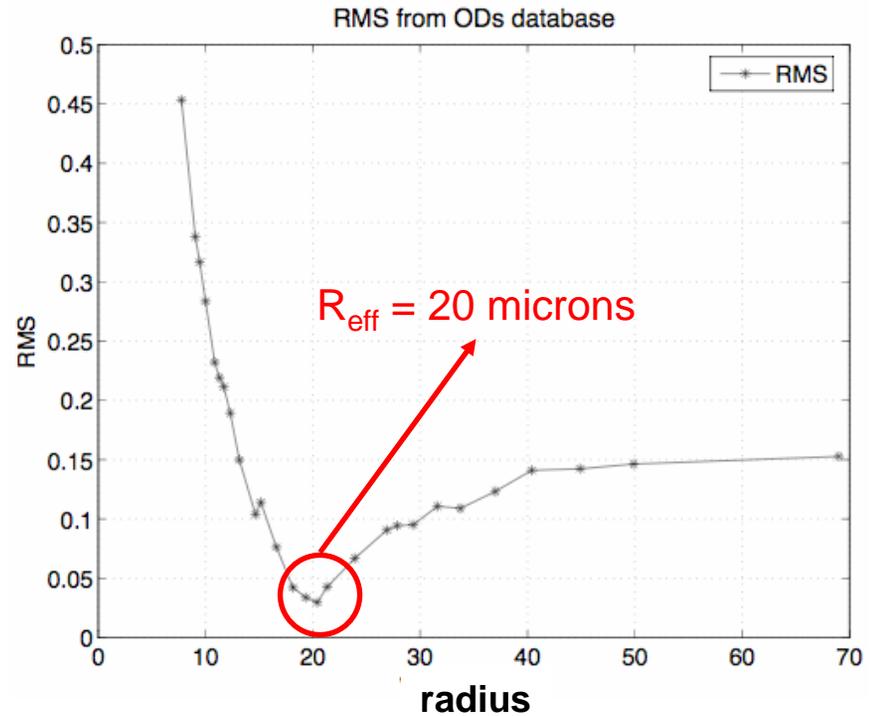
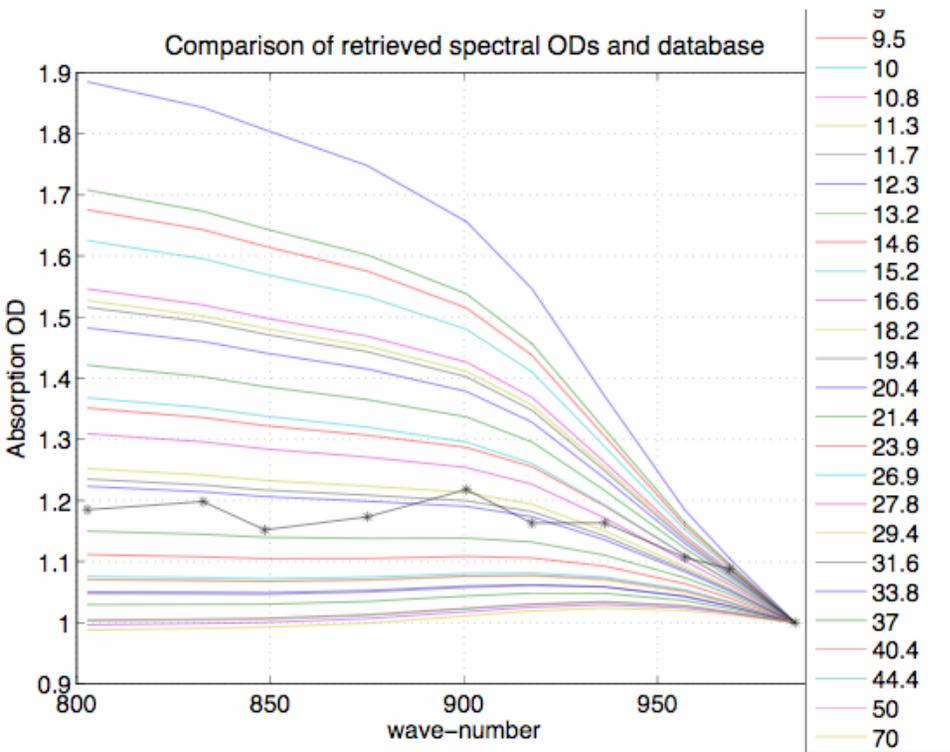
Simulation and retrieval: OD vertical non homogeneity



The real state: investigate the cloud “homogeneity”

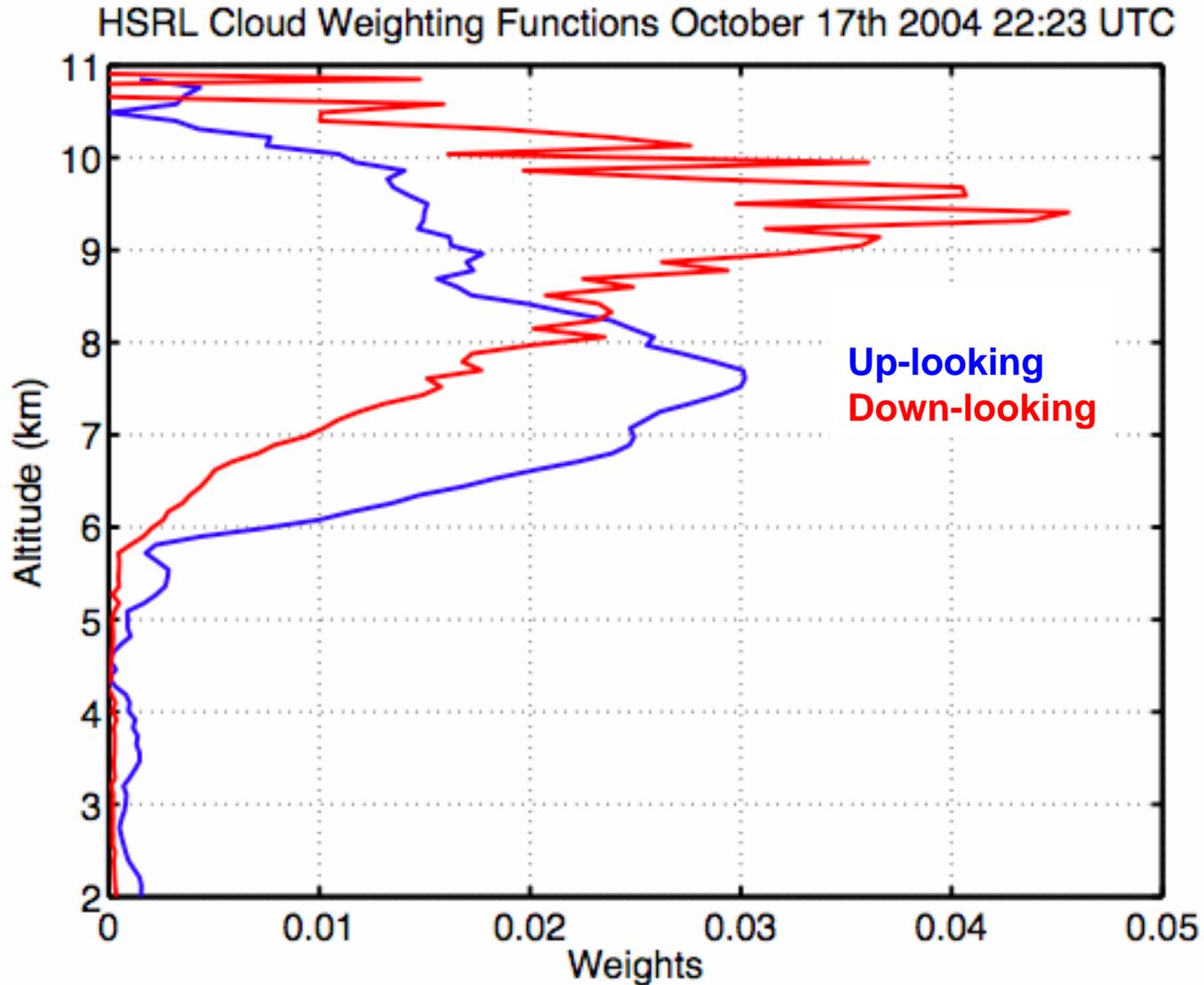


Retrieval of the effective dimension

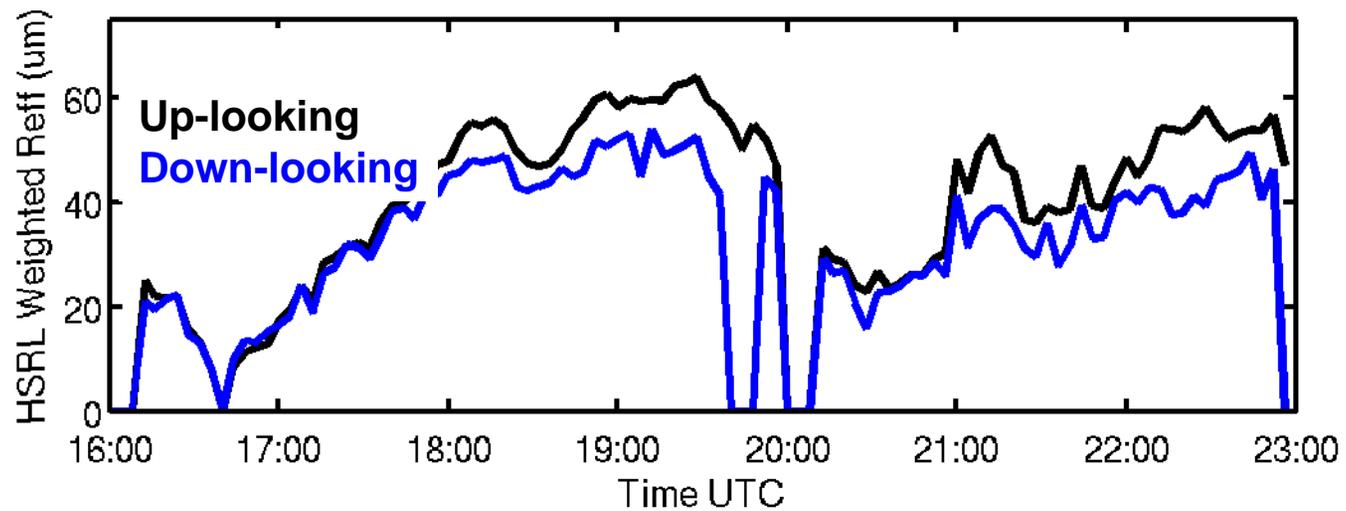
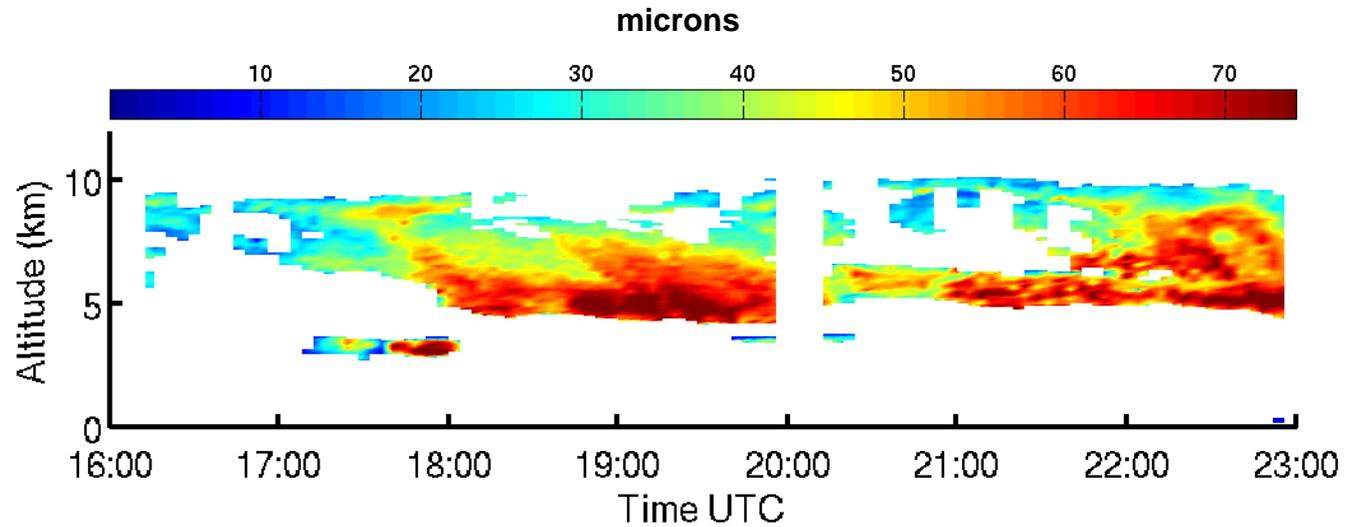


Lidar weighting functions

from layer to sensor transmittances



AHSRL/Radar effective radius retrieval

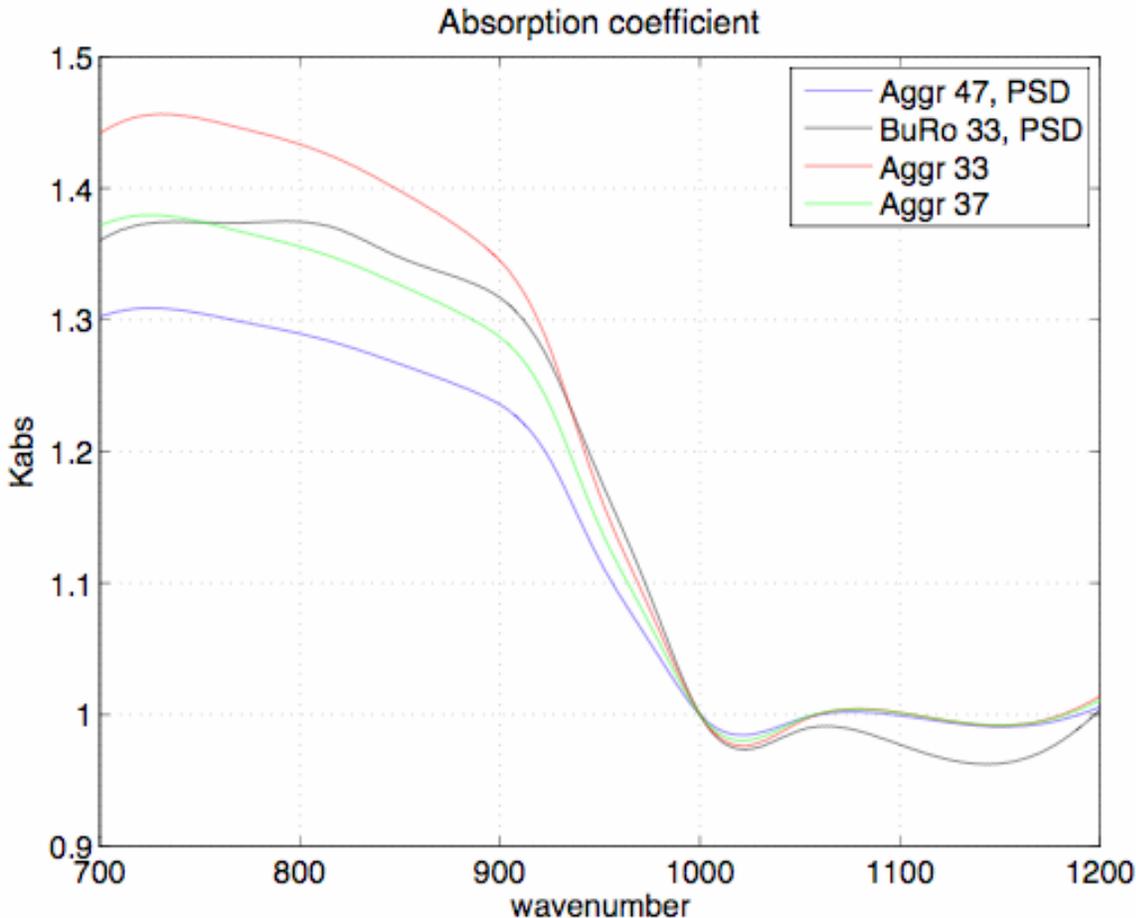


Retrieval of the effective dimension

AIRS

Bullett Rosettes: OD(900 cm^{-1})=1.3; OD(.532 μm)=1.46; R_{eff} =20 microns \rightarrow IWC=0.013 g/m^3
Aggregates: OD(900 cm^{-1})=1.3; OD(.532 μm)=1.36; R_{eff} =27 microns \rightarrow IWC=0.0097 g/m^3

[Mixture: OD=1.44, R_{eff} =33.5 microns *Xuebau Wu, Jun Li et al.* retrieval (droxtals, hexagonal and aggregates)]



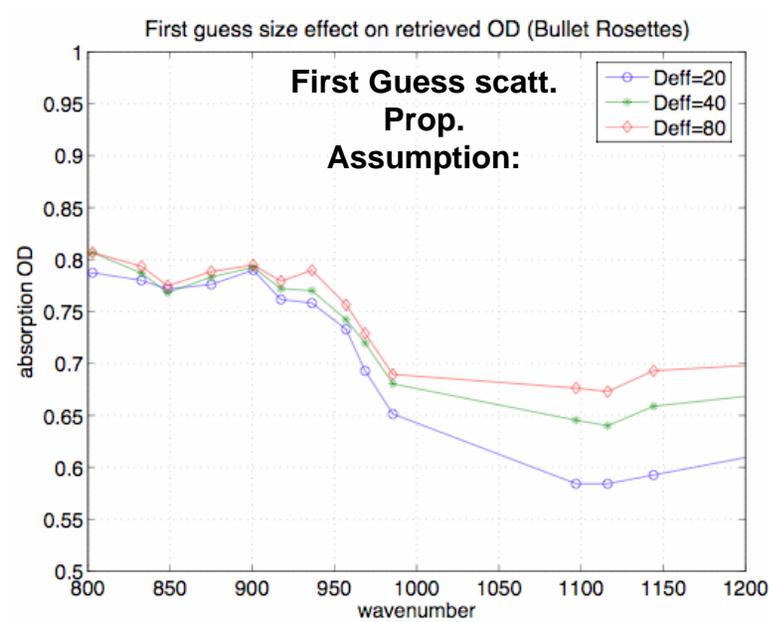
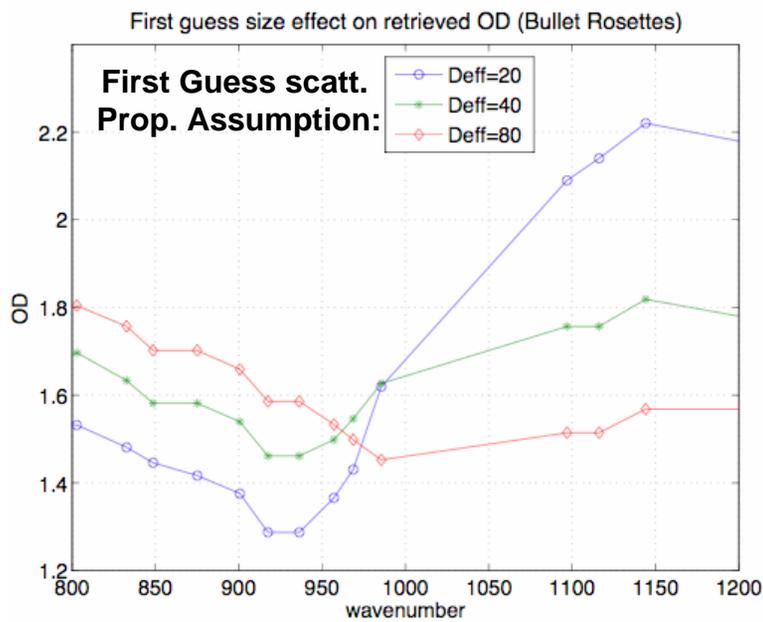
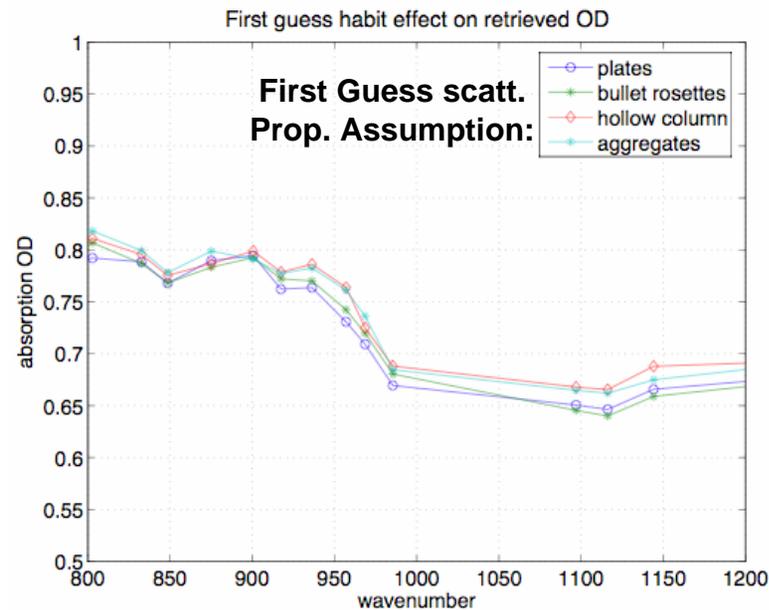
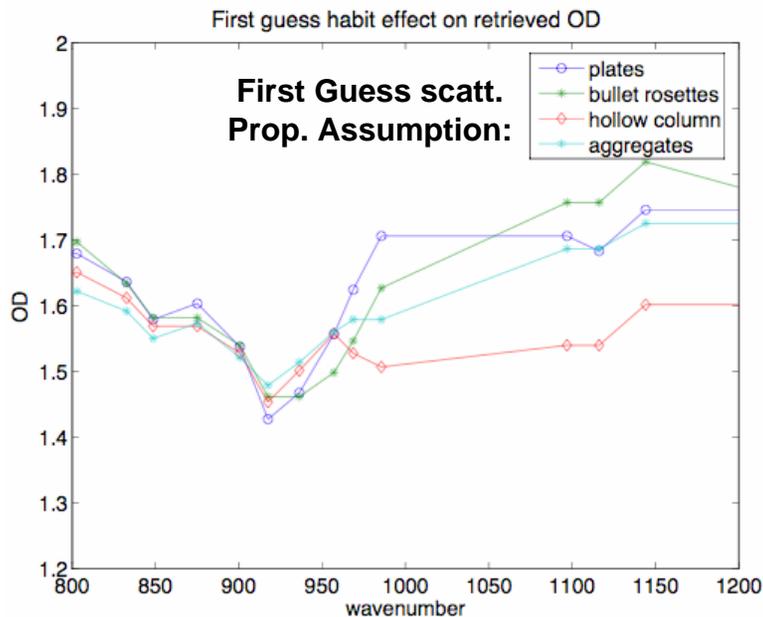
Considerations:

- IR optical properties effects on radiance are mostly defined by the particles' cross sectional area
- The effective dimension is volume/area dependent
- 'Bulky' particles have larger D_{eff} for the same PSD
- Comparison among different instruments must account for the habits and the PSDs assumptions
- AHSRL retrieves

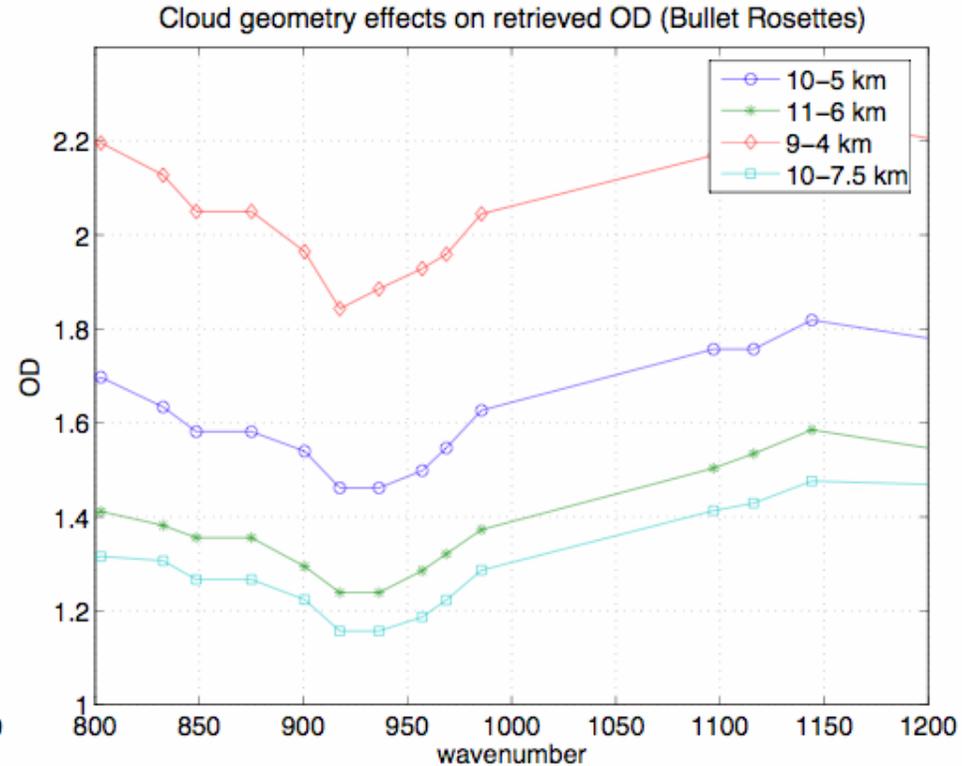
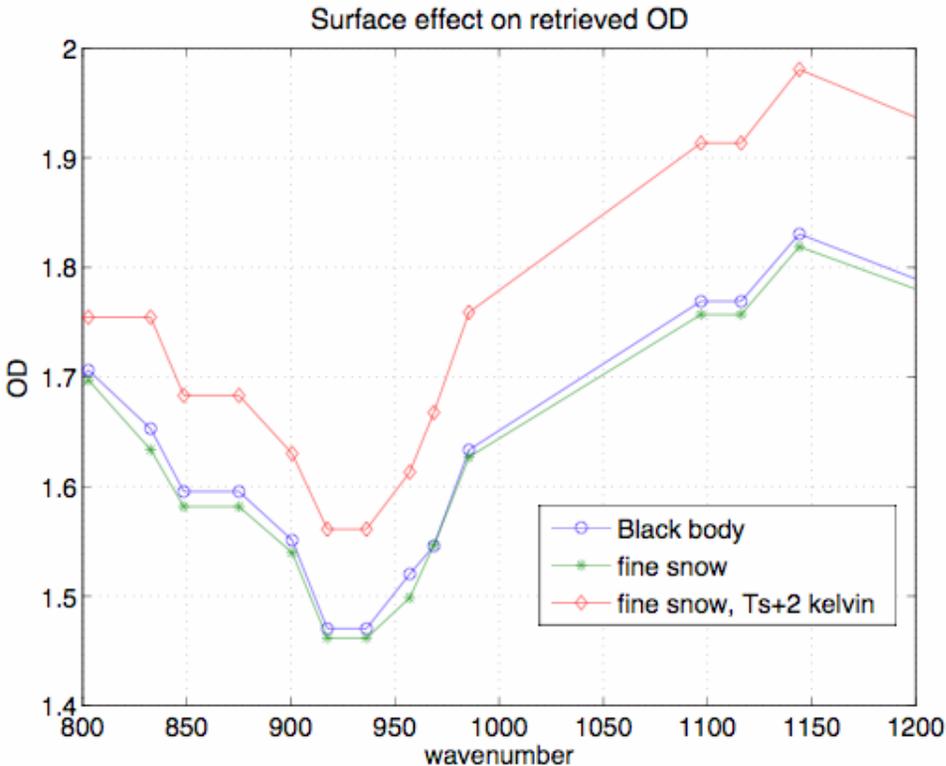
$$D'_{\text{eff}} = \sqrt[4]{\frac{9\langle V^2 \rangle}{\pi\langle A \rangle}}$$

without any assumption on habit and PSD
 \rightarrow Future comparison

The Algorithm sensitivity: Uncertainties on first guess assumption



The Algorithm sensitivity: Uncertainties on surface properties and cloud geometry



$$R \approx t_c \cdot (\varepsilon_s F_s) + (1 - t_c) \cdot F_c$$

Interpretation requires accounting for cloud emission and attenuation
 Ex. blu-cyan: cyan emits at lower T → has to attenuate less to get the same R

Conclusions

A new retrieval code working in the OD space and accounting for MS processes. RT-RET works with ground-airborne-satellite sensors

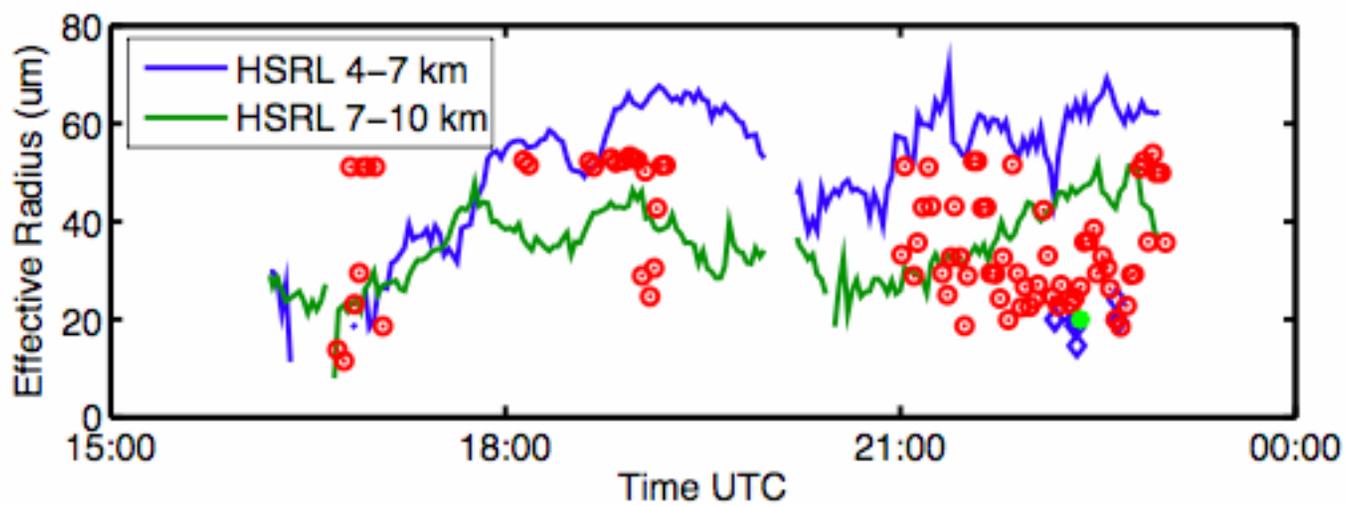
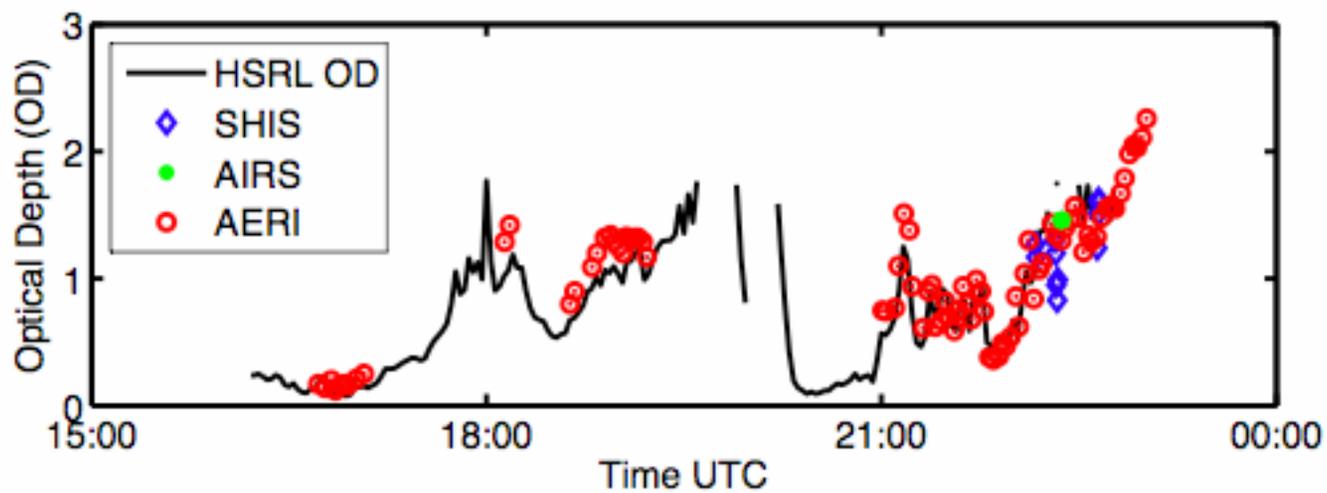
Limitations in cloud properties retrievals using IR only

Future

Improve in the speed/database/algorithm.

Validation case studies for certain geographical sites.

IR-Lidar/Radar consistent comparisons on particles retrieved dimensions: D'_{eff}

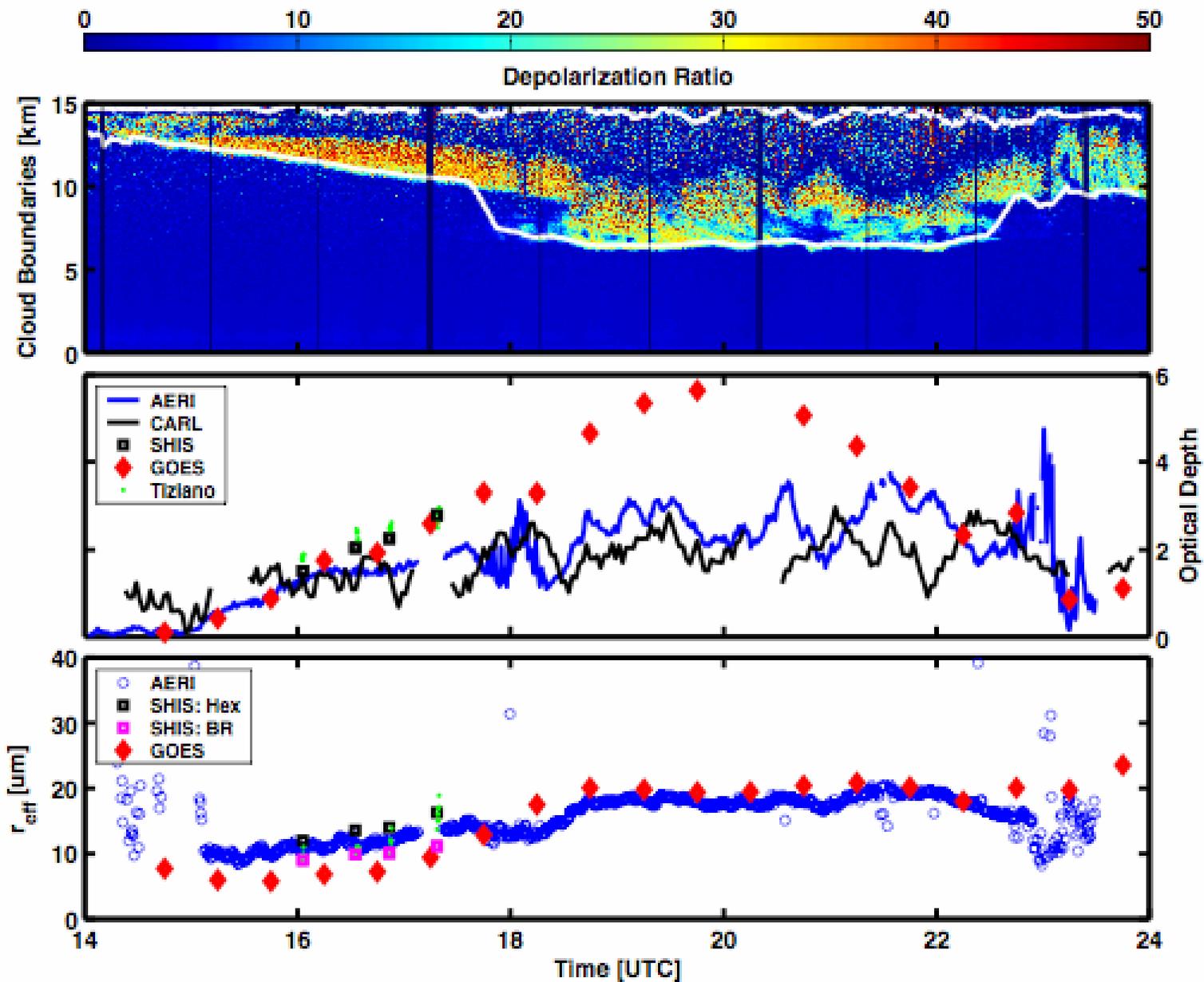


Effective Diameter Prime

is the fundamental quantity derived from the combination of lidar and radar backscatter cross sections, where $\langle V^2 \rangle$ and $\langle A \rangle$ refer to the average volume-squared and average area of the cloud particles as defined by Donovan and Lammeren JGR, v106, D21, pp 27425. The only user supplied quantity required for this computation is the backscatter phase function for ice crystals.

Could be computed from hsrl measurements for each point within the cloud, however the measurement is sensitive altitude averaging lengths and averaging times. Short averages are likely to have excessive noise. Long averages are likely to mix regions of low and high scattering. As a result, we have decided to allow the user to specify this value as a cloud average. It is recommended that an independent analysis of the data be used to determine an appropriate value for `p180_ice`. Errors in `effective_diameter_prime` are proportional to the 4th-root of the error in `.f`

Texas 2002



PARAMETERS DEFINING THE GAMMA

Gamma Type SD have been fitted to measured PSDs using:

1st, 2nd and 6th moments of the PSDs

Correlation coefficients $r^2 > 0.8$

$$N(D) = N_0 D^\mu e^{-\lambda D}$$

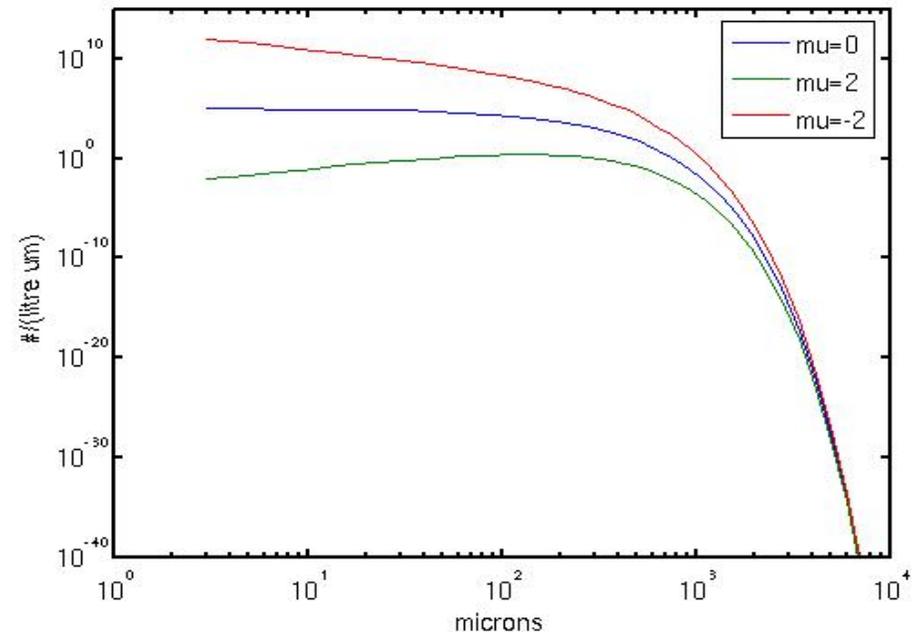
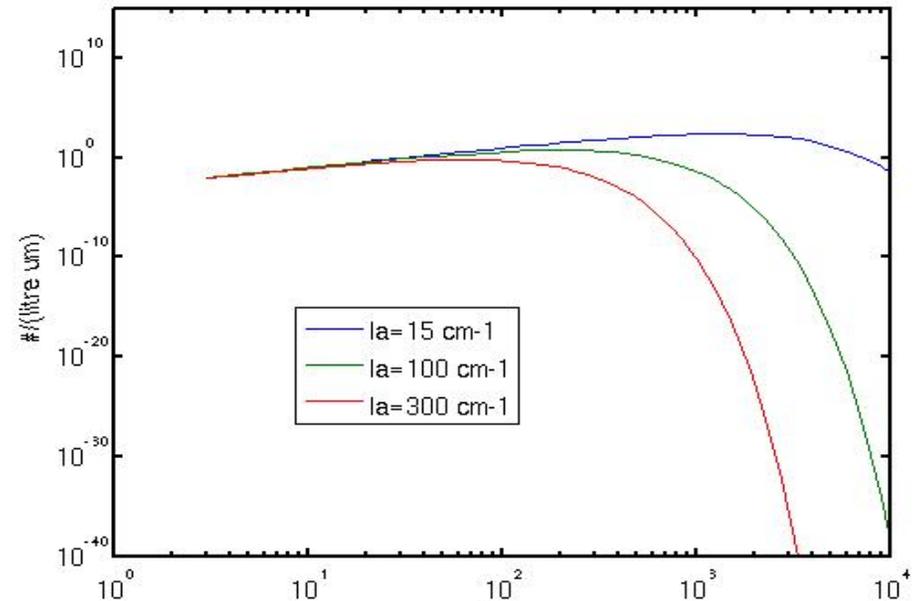
D = dimension (cm)

N_0 = intercept

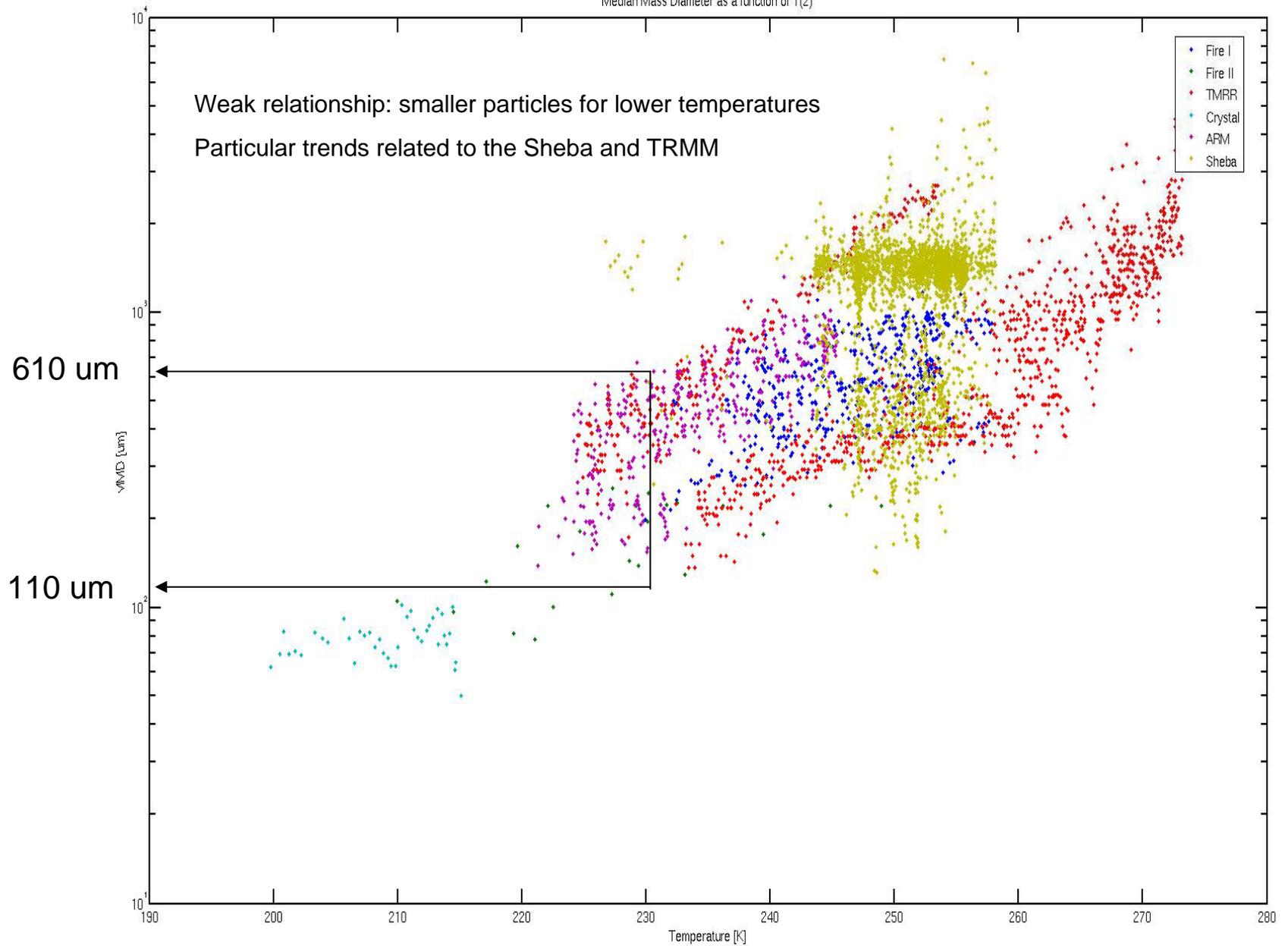
λ = slope

μ = dispersion (when =0 --> exp.)

Additional: IWC, DMM, T

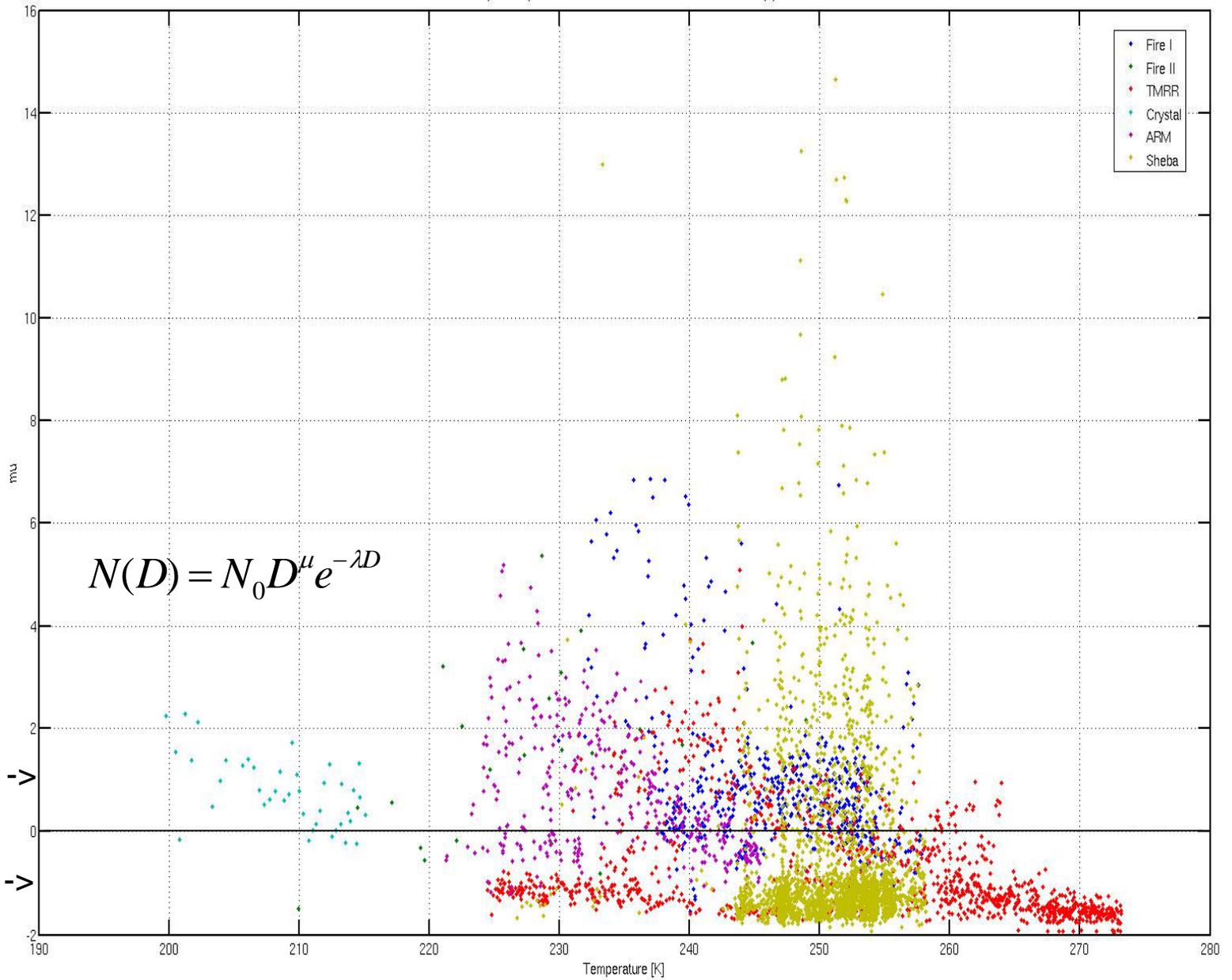


Median Mass Diameter as a function of T(z)



Ice clouds ←  → water/mixed phase?

Dispersion parameter of the Gamma SD as a function of T(z)



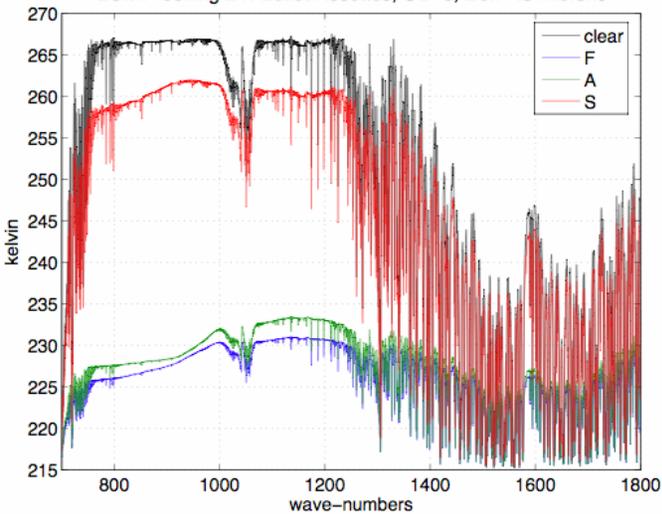
$$N(D) = N_0 D^\mu e^{-\lambda D}$$

Under exp ->

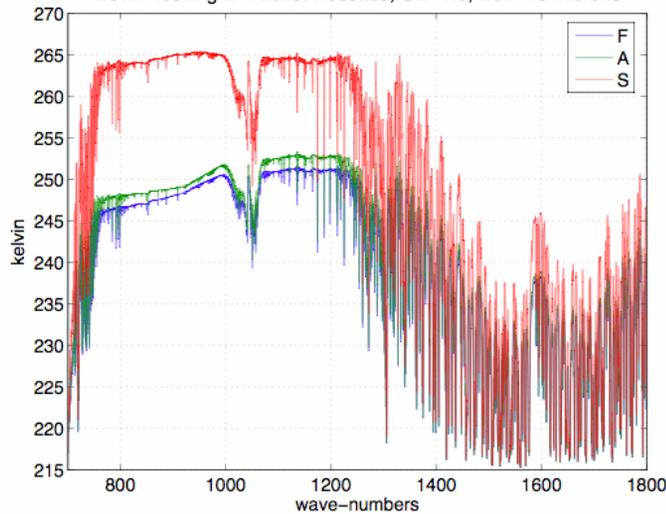
Over exp ->

TRMM deep convection --> over exponential

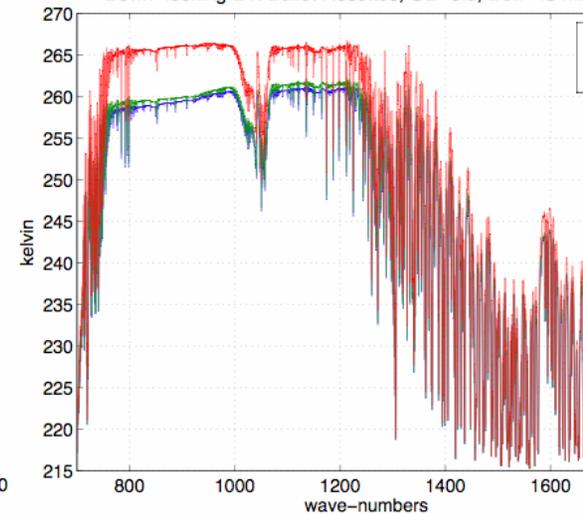
Down-looking BT: Bullet Rosettes, OD=5, Deff=43 microns



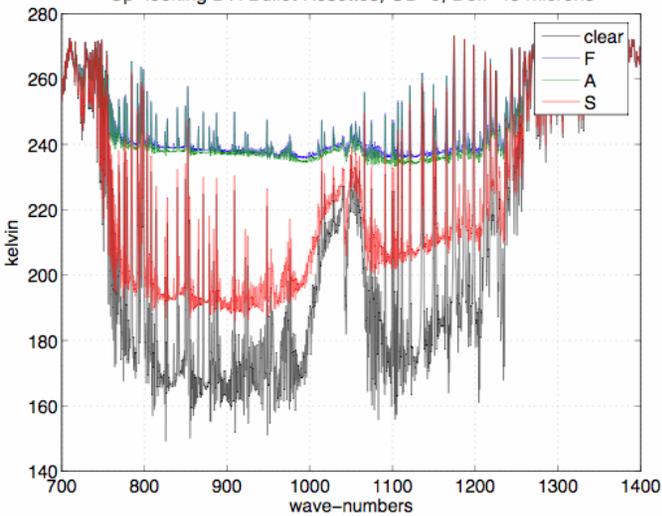
Down-looking BT: Bullet Rosettes, OD=1.5, Deff=43 microns



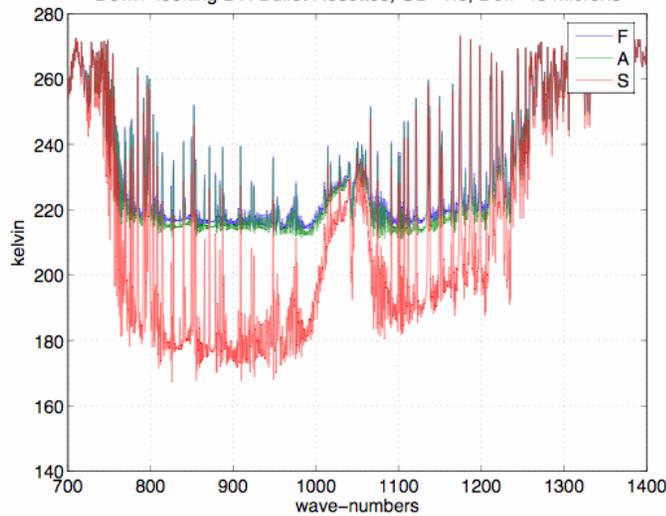
Down-looking BT: Bullet Rosettes, OD=0.5, Deff=43 microns



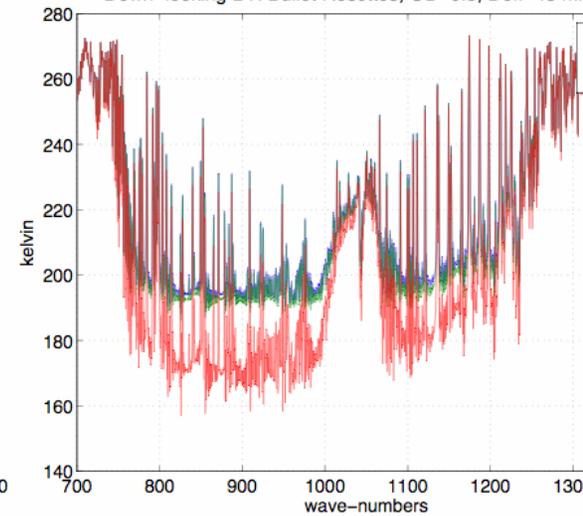
Up-looking BT: Bullet Rosettes, OD=5, Deff=43 microns

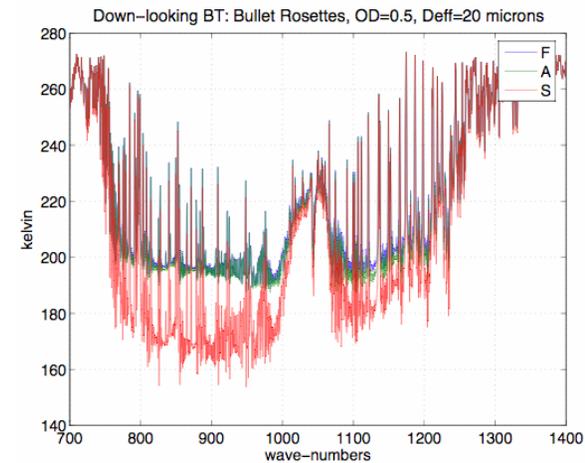
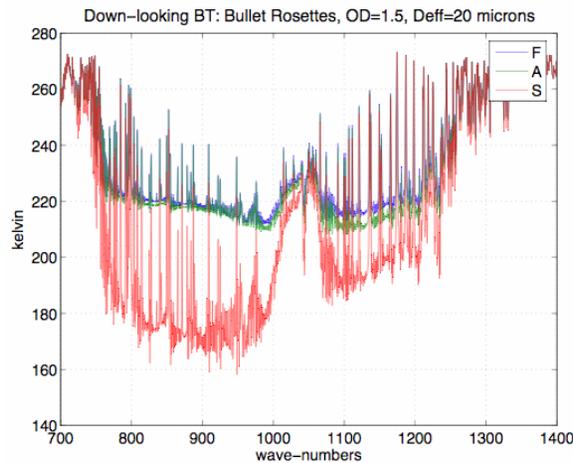
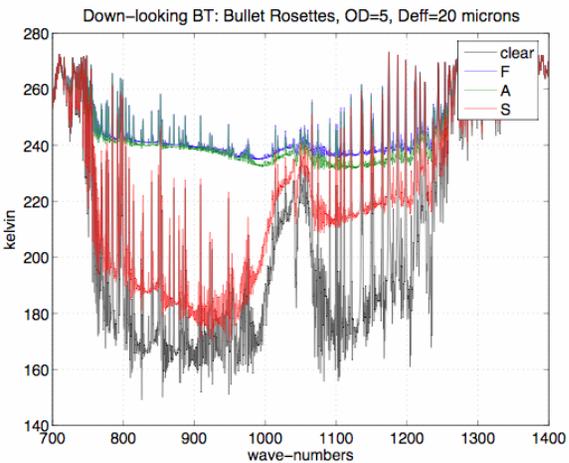
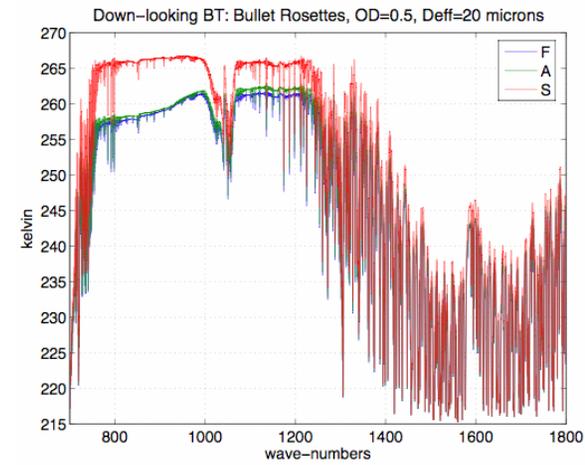
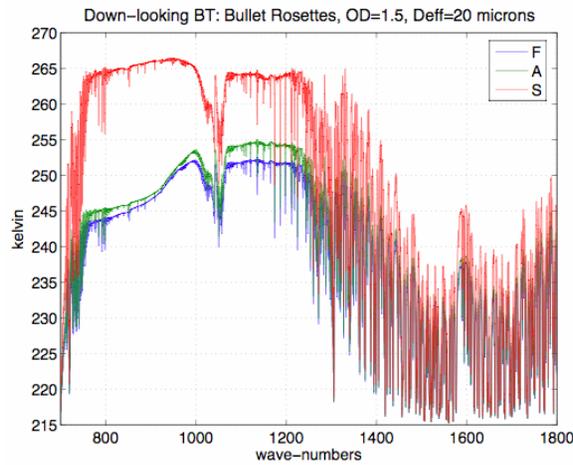
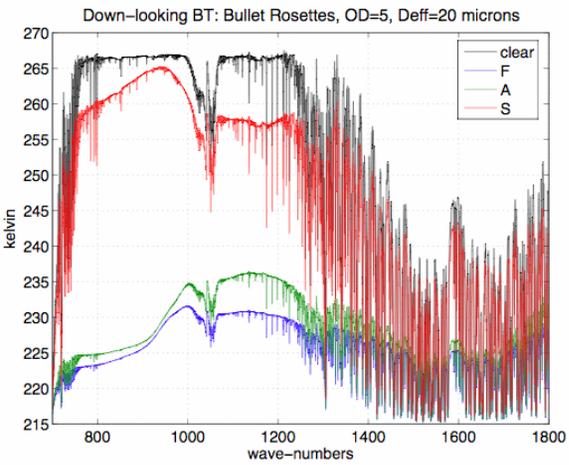


Down-looking BT: Bullet Rosettes, OD=1.5, Deff=43 microns



Down-looking BT: Bullet Rosettes, OD=0.5, Deff=43 microns





Up-looking Retrieval and AHSRL comparison

